

Late Miocene to late Pleistocene geomagnetic secular variation at high northern latitudes

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SUMMARY

We report a palaeomagnetic study of Icelandic lavas of late Miocene to late Pliocene age to test the geocentric axial dipole hypothesis at high northern latitudes. Cores were sampled from 125 sites in the Fljótisdalur valley in eastern Iceland, and hand samples were taken for 17 new incremental heating $^{40}\text{Ar}/^{39}\text{Ar}$ age determinations. 96 per cent of the cores were oriented using both a Brunton compass and a sun compass. Comparison of the magnetic and sun azimuths reveals deviations of $\pm 5^\circ$, $\pm 10^\circ$ and $\pm 20^\circ$, respectively, for 42, 16 and 3 per cent of the data points, indicating that core sampling intended for palaeosecular variation (PSV) studies at high northern latitudes should be oriented by sun. A total of 1279 independent specimens were subjected to AF- and thermal-demagnetization for palaeodirectional analysis, and well-grouped site mean directions were obtained for 123 sites of which 113 were found to be independent sites. Applying a selection criteria of $k > 50$ and $N \geq 5$ ($N_{\text{mean}} = 9.5$), we obtain a combined grand mean direction for 46 normal and 53 reverse (for $\text{VGP}_{\text{lat}} > \pm 45^\circ$) polarity sites of declination = 5.6° and inclination = 77.5° that is not significantly different from that expected from a GAD field. The corresponding palaeomagnetic pole position ($\text{VGP}_{\text{lat}} = 86.3^\circ\text{N}$, $\text{VGP}_{\text{lon}} = 21.2^\circ\text{E}$, $dp/dm = 4.0^\circ/4.3^\circ$) is coincident with the North Pole within the 95 per cent confidence limits. An updated age model is constructed based on the $^{40}\text{Ar}/^{39}\text{Ar}$ ages, showing that the majority of the Fljótisdalur lavas fall within 2–7 Ma. We combine the Fljótisdalur data with existing data from the nearby Jökuldalur valley. The 154 palaeodirections are well-dispersed between 1 and 7 Ma and constitute a high-quality data set for PSV analysis. Our results partly support previous conclusions of a generally higher dispersion during reverse polarity intervals. However, when comparing our Matuyama data with Brunhes age data from Jan Mayen, we find no evidence of a higher VGP scatter during the Matuyama as previously suggested. When comparing our VGP scatter to the two commonly used models for VGP dispersion: Model G and TK03, we find a good fit for all 1–7 Ma VGP scatter data $S_{B(1-7)}$ to Model G, whereas $S_{B(1-7)}$ is not fitted by TK03, even when considering the uncertainty of $S_{B(1-7)}$. We also find that all VGP scatter estimates, except that for the Gilbert subset, are consistent with Model G, while the discrepancy with TK03 is generally larger.

Key words: Thermochronology; Arctic region; Palaeomagnetic secular variation; Palaeomagnetism; Reversals: process, time scale, magnetostratigraphy.

1 INTRODUCTION

Satellite and ground observatory observations of the Earth's magnetic field (the geomagnetic field) show that it changes over time in both direction and intensity. These temporal changes are known

as secular variation and are thought to reflect both short- and long-term changes in fluid processes in the Earth's core (Johnson & McFadden 2007). A fundamental assumption of palaeomagnetism that the secular variation when observed at geological timescales, that is, palaeosecular variation (PSV), is averaged out and the

time-averaged geomagnetic field (TAF) is a geocentric axial dipole field: this hypothesis is known as the geocentric axial dipole (GAD) hypothesis (Hospers 1954). However, the question remains as to how long must the Earth's field be averaged to be a GAD (Lawrence *et al.* 2009; Aubert *et al.* 2010; Tauxe 2010; Korte *et al.* 2011; Cromwell *et al.* 2013, 2018; Pavón-Carrasco *et al.* 2014; Ziegler & Constable 2015).

By examining the temporal and spatial variations in the Earth's magnetic field over time using palaeomagnetic data, we may be able to test the validity of the GAD hypothesis. However, in order to conduct sound statistical studies of the TAF and PSV, we need global palaeomagnetic databases that are unbiased in terms of the temporal and spatial sampling for the period investigated with data being of high quality (Johnson *et al.* 2008).

In order to conduct statistical studies of the TAF and PSV, it has been argued that an ideal data set would sample the geomagnetic field on the order of 10^6 years from about 100–200 sites (Ziegler *et al.* 2011; Lhuillier & Gilder 2013). Cromwell *et al.* (2018) have further considered the minimum number of ChRMs per site needed to confidently trust individual palaeodirections as input for TAF and PSV studies. They suggest a data selection criteria with $N \geq 4$ (minimum number of ChRMs per site) and $k \geq 50$. This cut-off is based on simulations by Cromwell *et al.* (2018) that indicate that within-site estimates of the circular standard deviation (c_w) is reduced by ~ 35 per cent for $N = 4$ as compared to $N = 2$. Interestingly, the same study shows that the variance of c_w is reduced by further ~ 27 per cent for $N = 10$ as compared to $N = 4$, while an improvement of only ~ 7 per cent is obtained for $N = 15$ as compared to $N = 10$. This suggests that it is worthwhile obtaining data sets with $N = 10$,

Early attempts to compile palaeomagnetic global data sets for modelling the PSV and TAF over geological timescales of 10^5 – 10^6 yr indicated that many of the older palaeomagnetic results are of too low quality to truly test the GAD hypothesis (Johnson & Constable 1997). Modern palaeomagnetic data sets with relatively strict data standards have since then been collected and used as input for global palaeomagnetic databases to model, for example, the 0–5 Ma geomagnetic field (Johnson *et al.* 2008) and recently also the 0–10 Ma field (Cromwell *et al.* 2018). However, considerable temporal and geographical biases remain in the databases with a majority of data points being less than 500 kyr old and collected at low- to mid-latitudes (Johnson *et al.* 2008; Cromwell *et al.* 2018).

The basaltic lava piles of the Iceland plateau, exposed onshore in deeply dissected valley and fjord systems (64 – 66°N), are well suited for palaeomagnetic studies due to their fidelity at recording primary remanent magnetizations, the large number of accessible flows with minimal hydrothermal alteration covering the past 15–16 Myr (Hardarson *et al.* 1997), and the limited post-emplacement tectonic disturbance (Kristjánsson & Jonsson 2007). Despite a large collection of palaeomagnetic studies over the past 50–60 years (McDougall *et al.* 1977; Watkins & Walker 1977; Kristjánsson & Jóhannesson 1999; Kristjánsson 2013), only few studies report TAF and PSV behaviour for rocks older than 0.5 Ma using modern-standard quality criteria (Johnson *et al.* 2008; Cromwell *et al.* 2018). Instead, a large majority of the studies focus on either transitional polarity events and excursions or on detailed magnetostratigraphic correlations (Dagley *et al.* 1967; Watkins & Walker 1977; Kristjánsson 2008; Camps *et al.* 2011; Jicha *et al.* 2011), with the latter studies mostly being based on two samples per flow. A compilation study of 0–16 Ma Icelandic lavas was presented by Kristjánsson (2013), covering remanence results of more than 5200 lava flows that were collected over the past 50 years. Despite that the majority of the

studies are based on two samples per site, an interesting outcome of the study is an apparently smooth and continuous frequency distribution of Icelandic virtual geomagnetic poles (VGPs) in latitude, suggesting that secular variation, excursions and transitions could be treated jointly in VGP statistical studies as manifestations of turbulent convective processes in the Earth's core.

In 2013 and 2014, we visited the large Fljótisdalur valley in eastern Iceland (Fig. 1) with the purpose of providing high-quality palaeodirectional, palaeointensity and $^{40}\text{Ar}/^{39}\text{Ar}$ data for the late Miocene to late Pliocene period from high northern latitudes. The valley was originally sampled by Dagley *et al.* (1967) and Watkins & Walker (1977), collecting two cores per site. In total, we sampled 125 sites along five sections in Fljótisdalur as part of an eight weeks sampling campaign that also included sampling of late Pliocene–Pleistocene rocks in the neighbouring Jökuldalur valley (Fig. 1). Palaeodirectional and palaeointensity data from 45 distinct lavas flows, including 11 new $^{40}\text{Ar}/^{39}\text{Ar}$ ages, were published from Jökuldalur (Døssing *et al.* 2016).

Here, we present an analysis of the palaeodirectional data from Fljótisdalur, including a revised age model of the lava flows in the valley based on 17 new $^{40}\text{Ar}/^{39}\text{Ar}$ age determinations. The age of the lavas in Fljótisdalur allows us to combine the results in this study with the results of Jökuldalur in order to construct an estimate of the PSV for high northern latitudes for the late Miocene to Pleistocene. Hence, we present high-quality estimates of PSV for the 1–7 Ma time interval by making a joint analysis of the Fljótisdalur and Jökuldalur palaeodirectional data.

2 GEOLOGY AND PREVIOUS WORK

Iceland is located in the Northeast Atlantic at 63 – 67°N . The island itself sits at the junction between the North Atlantic Ridge and the Greenland–Iceland–Faroe Islands ridge and atop anomalously hot mantle. In this unique extensional tectonic setting, the active volcanic rift zones (see inset map in Fig. 1), that is, the northern volcanic zone (NVZ), the eastern volcanic zone (EVZ) and the western volcanic zone (WVZ), are infilled with basaltic extrusive and intrusive rocks. During this progressive accumulation of volcanic material the areas around the volcanic zones cool, contract and subside continuously, causing the lava sequences to tilt gently towards the axes of the volcanic zones (Pálmason & Sæmundsson 1974).

The southwest–northeast striking Fljótisdalur valley is located on the Eurasian Plate in eastern Iceland, about 80 km east of the NVZ (Fig. 1). A comprehensive geological mapping of the valley was carried out in the late 1970s as part of an M.Sc. project by Gudmundsson (1978). However, the first palaeomagnetic campaign in Fljótisdalur, collecting two cores per site, was already conducted in 1964 and 1965 as part of a large sampling campaign that focused on regional magnetostratigraphic correlation across eastern Iceland (Dagley *et al.* 1967; Watkins & Walker 1977). Based on this regional stratigraphic framework, Dagley *et al.* (1967) and Watkins & Walker (1977) suggested that the lavas in the upper southwestern part of Fljótisdalur, that is, along sections FR to FV, are of late Miocene to early Pleistocene age. This was later supported by six $^{40}\text{Ar}/^{39}\text{Ar}$ ages (Mussett *et al.* 1980). Watkins & Walker (1977) also found that the lava pile in eastern Iceland is characterized by an overall gently westerly dip and that the dip generally increases linearly with depth (and hence age) below the original surface. This conclusion was based on several hundred dip angle measurements of the lavas across eastern Iceland, including Fljótisdalur.

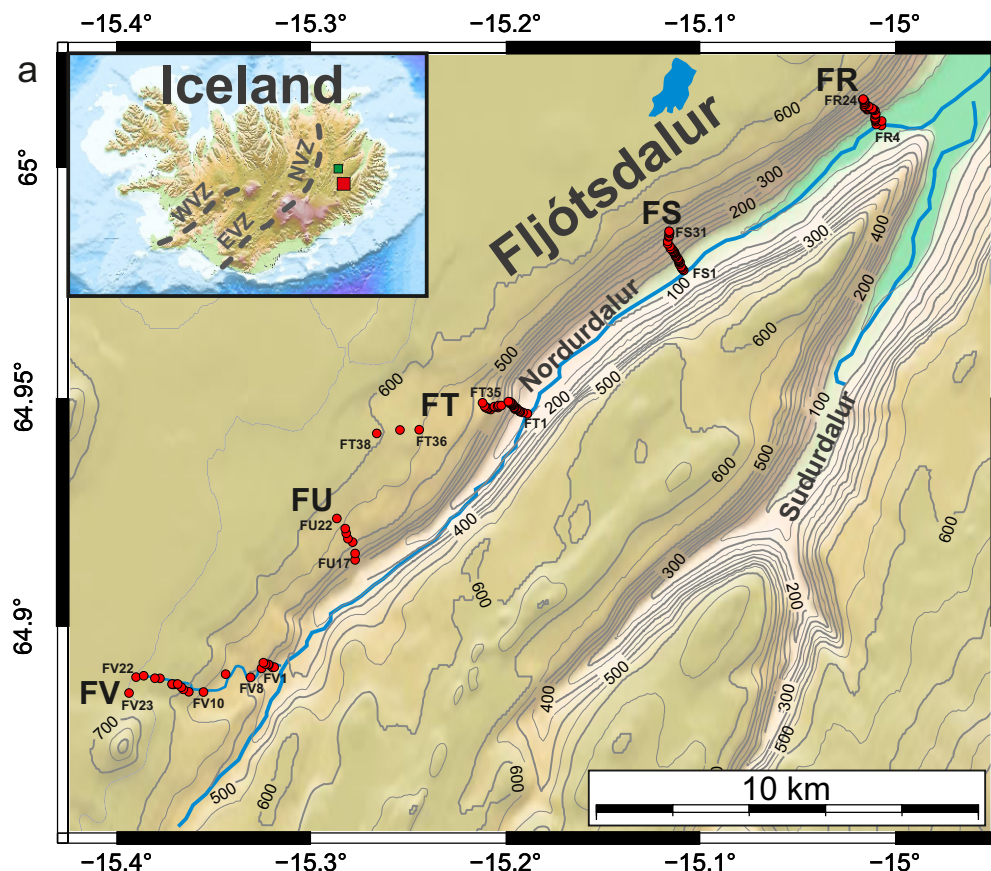


Figure 1. Topographic map of the upper part of the Fljótsdalur valley. ‘Fljótsdalur’ here refers to the main tributary valley Nordurdalur. Red dots: Location of drilled sites along sections FR, FS, FT, FU and FV. Inset map of Iceland: red square, location of Fljótsdalur; green square: location of Jökuldalur where full-vector palaeomagnetic PSV results of 0.5–3.1 Ma rocks have recently been published (Døssing *et al.* 2016). Black dashed lines: main volcanic rift zones. Abbreviations: EVZ, eastern volcanic zone; NVZ, northern volcanic zone; WVZ, western volcanic zone.

In this study, we have revisited the basaltic lavas along sections FR to FV in Fljótsdalur. The five sections were initially mapped by Dagley *et al.* (1967) and Watkins & Walker (1977), who named the five sections R, S, T, U and V. The lavas along the sections are ideal for PSV studies because they are well-mapped, easily accessible and exhibit a westerly dip less than 10° (Watkins & Walker 1977), because they are located outside a regional volcanic flexure zone to the northeast.

3 METHODS

3.1 Core sampling

We collected >1500 individual cores from 125 sites along sections FR, FS, FT, FU and FV (Figs 1 and 2) for palaeodirectional and palaeointensity analysis. The results of the palaeointensity analysis will be presented in a future study. Our core sampling was restricted to unaltered massive cores of basaltic lavas in well-defined flow units, and cores were preferably drilled over tens of metres to avoid the potential influence of lightning strikes and/or non *in situ* blocks. In compound flow units, we selected only one subunit for coring (e.g. FR8, FR19, FT11, FT17, etc.).

All cores were oriented using a Brunton magnetic compass. In addition, 96 per cent of all cores were oriented using a sun compass. This provides us with a unique data set for comparing magnetic and sun azimuths at high northern latitudes. As observed in Fig. 3,

we find that 42, 16 and 3 per cent of the magnetic azimuth estimates (corrected to true north) deviate more than $\pm 5^\circ$, $\pm 10^\circ$ and $\pm 20^\circ$, respectively, from the corresponding sun azimuth estimates. These numbers strongly support the conclusion of Cromwell *et al.* (2013) that core sampling intended for high-quality PSV studies at high northern latitudes should preferably be oriented by sun (or alternatively by a differential GPS).

3.2 $^{40}\text{Ar}/^{39}\text{Ar}$ radiometric dating

Hand samples of fresh rock were collected from all sites to deliver potential new $^{40}\text{Ar}/^{39}\text{Ar}$ radiometric age determinations. We selected 17 basalt samples for $^{40}\text{Ar}/^{39}\text{Ar}$ age determination based on their response to incremental heating age dating and the need for improved stratigraphic control. The experiments were carried out at the Argon Geochronology Laboratory at Oregon State University. A detailed description of the sample preparation and dating method is found in Supporting Information Section S1-1.

3.3 Demagnetization experiments

All palaeomagnetic laboratory work was carried out at the palaeomagnetic laboratories of Imperial College London (ICL) and University of Oxford. Upon cutting the samples, the natural remanent magnetization (NRM) was measured using an AGICO JR5A spinner magnetometer and a Molspin Minispin at ICL or a Molspin and

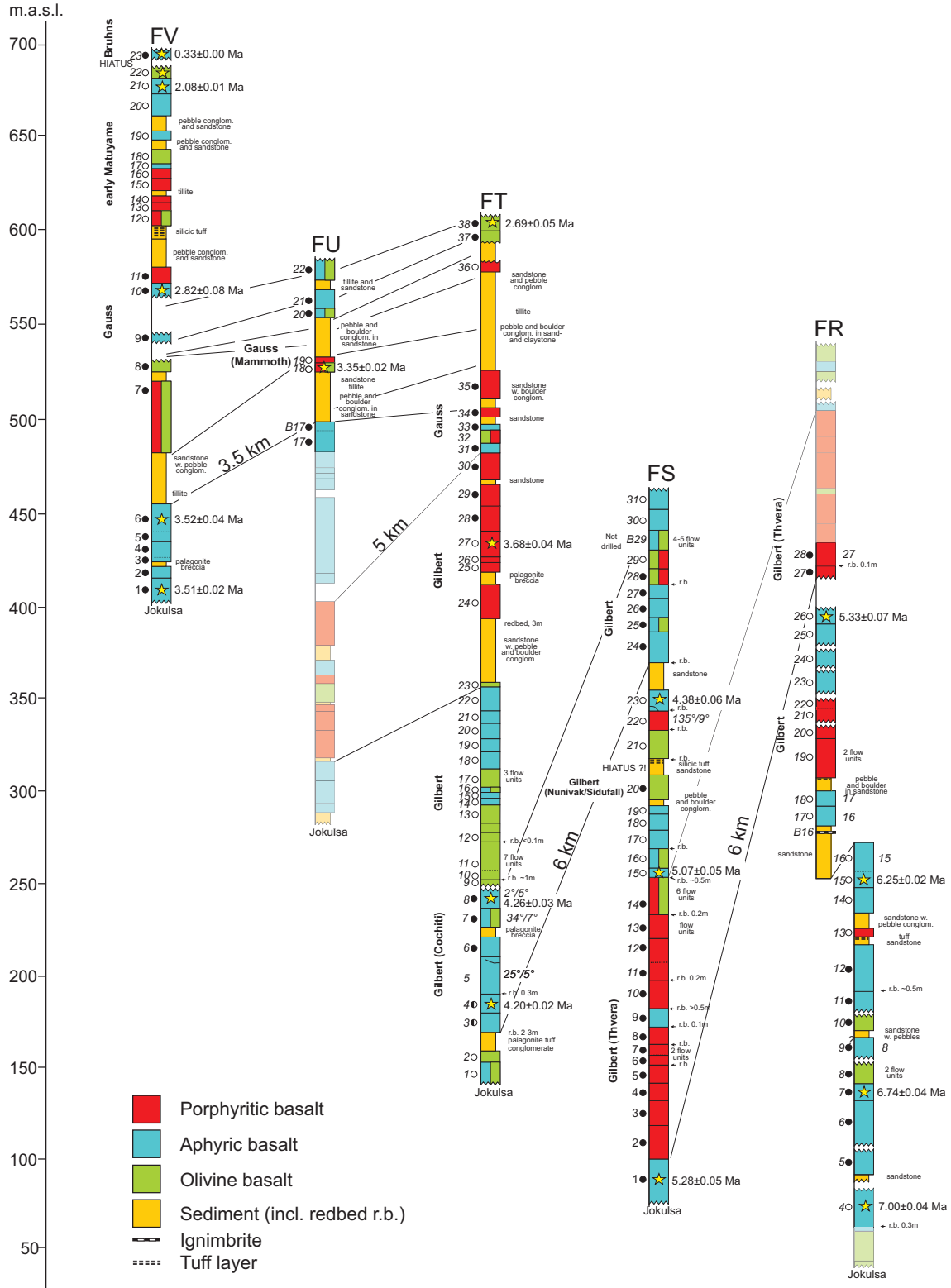


Figure 2. Updated magnetostratigraphy of sections FR to FV in Fljótsdalur based on the results presented here. Stratigraphic correlation of the sections is based on visual inspection in the field, site type geology as well as new palaeomagnetic directions and $^{40}\text{Ar}/^{39}\text{Ar}$ radiometric ages (this study). Labels to the left are site labels (this study). White(black) filled circles are magnetic reverse(normal) polarity based on virtual geomagnetic pole (VGP) polarity (this study). Labels to the right are Ar–Ar ages in Ma (this study) and on-site strike/dip measurements where possible (dip to the west). The updated stratigraphic correlation builds upon the original stratigraphy of Gudmundsson (1978). Please note that the base and top flows in section FR were not drilled due to time constraints and redundancy with section FS. Similarly, we only drilled the top flows in section FU due to redundancy of the base flows of section FU with flows in section FV.

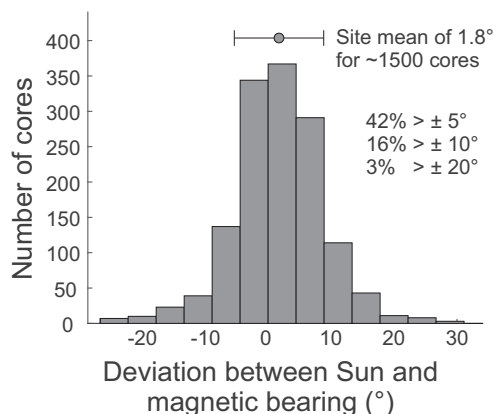


Figure 3. Histogram of difference between Sun and magnetic azimuth estimates for the ~ 1500 cores in this study.

three-axis 2G superconducting magnetometer at University of Oxford. Stepwise alternating-field (AF)- and thermal-demagnetization behaviours were measured on 1279 independent specimens from 125 sites for palaeodirectional analysis. Based on our results from Jökuldalur (Døssing *et al.* 2016), we subjected 1–2 specimens per site to thermal demagnetization using a magnetic measurements thermal demagnetiser (Oxford). The remainder of the specimens per site were subjected to AF magnetic cleaning using a Molspin Tumbling AF demagnetizer and a DTECH D-2000 high-performance AF demagnetizer at ICL, or a 2G magnetometer fitted with in-line AF-demagnetization coils at University of Oxford. Thermal- and AF-demagnetization experiments were carried out in 10–12 steps with maximum unblocking temperatures normally around 590 °C and within applied field magnitudes of 5–200 mT at ICL or 5–140 mT at Oxford. The 10–12 demagnetization steps were conducted to ensure proper removal of any viscous remanent magnetization and optimal determination of the characteristic remanent magnetization (ChRM) from principal component analysis (Kirschvink 1980).

4 RESULTS

4.1 $^{40}\text{Ar}/^{39}\text{Ar}$ radiometric ages

The 17 new high-resolution incremental heating $^{40}\text{Ar}/^{39}\text{Ar}$ age determinations are summarized in Table 1 and the age spectra are presented in Fig. 4. Assessments of each age-determination experiment are given in Supporting Information Section S1-2 along with inverse isochron diagrams (Supporting Information Fig. S1-1). The ArArCALC age calculation files will be made available for download from the EarthRef.org Digital Archive (ERDA).

In summary, the 17 new $^{40}\text{Ar}/^{39}\text{Ar}$ ages indicate that sections FR(FR4) to FV(FV23) span the interval of 7.00 ± 0.04 Ma to 334.4 ± 4.2 Ka (Table 1). A significant hiatus of 1.75 Myr is found between in the uppermost part of section FV between sites FV22 (2.08 ± 0.01 Ma) and FV23 (334.4 ± 4.2 Ka). Between 5 and 31 of the heating steps define the horizontal age plateaus that also include from 18 to 100 per cent of the $^{39}\text{Ar}_K$ gas released. Six of the 17 groundmass separates failed to yield true plateau ages (> 50 per cent of the $^{39}\text{Ar}_K$ gas released) and for those samples we use the term mini-plateau age (weighted mean age) (Fig. 4). Fourteen out of 17 inverse isochron intercepts are within error of the 295.5 modern atmospheric ratio of $^{40}\text{Ar}/^{36}\text{Ar}$ (Supporting Information Section S1 and Fig. S1-1). For all but two experiments the K/Ca value decreases from the low-temperature steps through the age

plateau to the high-temperature steps (Fig. 4); most likely a result from alteration phases that readily lose gas at low temperatures while argon retentive phases such as groundmass plagioclase and clinopyroxene tend to degas at high temperatures. Our approach resulted in high precision ages (with nominally 0.3–2.7 per cent 2σ uncertainties, including J -value errors).

The samples from sections FV, FU, FT and down to FS-15, with the exception of FV-6, generally yielded concordances in plateau, isochron and total fusion ages, indicating that the primary argon reservoirs were not significantly affected by alteration or recoil, and that the K–Ar isotope system has effectively remained closed since the time of eruption (Fig. 4). The data show only modestly disturbed age spectra (^{39}Ar recoil during irradiation), if disturbed at all. Sample FV-6 yielded a more disturbed age spectrum for which we defined a mini-plateau age of 3.52 ± 0.04 Ma (26 per cent $^{39}\text{Ar}_K$). An alternative interpretation is a slightly older plateau age of 3.72 ± 0.06 Ma (low-mid temperature steps, 56 per cent $^{39}\text{Ar}_K$) that is concordant with an inverse isochron age of 3.74 ± 0.10 Ma. However, this interpretation is at odds with the very concordant FV-1 age data of 3.51 ± 0.02 Ma (97 per cent $^{39}\text{Ar}_K$). Sample FS-1 and all samples from FR yielded disturbed age spectra, in particular FR-15, FR-7 and FR-4, hence rendering determination of the age of crystallization difficult. The groundmass material seems to be affected by both $^{39}\text{Ar}_K$ recoil during irradiation as well as argon loss likely caused by low-temperature alteration following burial in the lava pile. The latter is a common problem when attempting $^{40}\text{Ar}/^{39}\text{Ar}$ age determinations on material collected at low elevation from lava piles of Miocene age in Iceland. We define mini-plateau ages at lower-temperature steps for FR-15 (6.25 ± 0.02 Ma), FR-7 (6.74 ± 0.04 Ma) and FR-4 (7.00 ± 0.04 Ma), while alternative interpretations at higher-temperature steps with resultant younger ages (5.5–6.0 Ma for FR-7 and FR-4) can also be made. However, such young ages cannot be reconciled within the framework imposed by the magnetostratigraphy (Fig. 2) and absolute age constraints imposed by lava flows higher in the stratigraphy. Hence, we consider the reported plateau ages closer to the actual crystallization ages.

For comparison we have reassigned the ages for legacy $^{40}\text{Ar}/^{39}\text{Ar}$ age determinations reported by Ross & Mussett (1976) and Mussett *et al.* (1980) on samples from these sections (V18-1, 2.09 Ma; V8-3a, 2.58 Ma; T17A-2a, 3.53 Ma; S19-1a, 4.68 Ma; R20A-2a, 4.62 Ma; R16-2a, 5.62 Ma; R5-2a, 6.01 Ma) by correcting for their reported decay constant and monitor standard age values so that the recalculated legacy age data is now comparable to the age determinations reported in Table 1. The reassigned legacy ages reported for units FV18 and FV8 are broadly consistent with the ages that we report for this stratigraphic interval. Still, the reassigned age of FV8 is *c.* 200 kyr younger than the plateau age for FV10 (with total fusion age of 2.56 Ma). We argue that the reported age for FV8 was somewhat influenced by argon loss. The reassigned age for FT17 is *c.* 200 kyr younger than FT27 (above) and *c.* 700–800 kyr younger than FT8 (below). We regard the age of FT17 as unreliable and influenced by argon loss. The reassigned age for sample FS19 corroborates well with the ages that we report for FS15 and FS23. The reassigned ages for FR20, FR15 and FR5, however, are 0.7–1.0 Myr younger than the age range implied for this stratigraphic interval by the new age determinations of FR26, FR15, FR7 and FR4. The reassigned ages are also consistently younger than the total fusion ages for FR26, FR15, FR7 and FR4. We attribute the young ages for FR20, FR15 and FR5 to the effects of argon loss and inadequate leaching pre-treatment. Deeply buried low-potassium basalts are notoriously challenging to obtain reliable K/Ar and Ar/Ar ages from, which is also evident from the complex

Table 1. Incremental heating $^{40}\text{Ar}/^{39}\text{Ar}$ age determinations for basalts from Fljótisdalur. ID, experiment ID; gm, groundmass; n, the number of heating steps in the plateau; N, total number of heating steps for the entire experiment. K/Ca values are calculated as weighted means for the age spectra. MSWD values for the age plateaus and inverse isochrons are calculated using N-1 and N-2 degrees of freedom, respectively. All sample information from this study where monitored against FCT sanidine (28.201 ± 0.023 Ma, 1σ) as calibrated by Kuiper *et al.* (2008). Reported errors on the $^{40}\text{Ar}/^{39}\text{Ar}$ ages are on the 95 per cent confidence level (2σ) including 0.2–0.3 per cent standard deviation in the J -value. All ages were calculated using the corrected Steiger & Jäger (1977) decay constant of $5.530 \pm 0.097 \times 10^{-10} \text{ yr}^{-1}$ (2σ) as reported by Min *et al.* (2000). For all other constants used in the age calculations we refer to Koppers *et al.* (2003). Incremental heating plateau ages and isochron ages were calculated as weighted means with $1/\sigma^2$ as weighting factor (Taylor 1997) and as YORK2 least-squares fits with correlated errors (York 1968) using the ArArCALC v2.6.2 software from Koppers (2002) available from the <http://earthref.org/ArArCALC/> website. Abbreviations: Mat: Material; Ic: Intercept.

Sample inf			Age spectrum						Total fusion		Inverse isochron		
Site	ID	Mat	Age $\pm 2\sigma$ (Ma)	^{39}Ar (%)	K/Ca	MSWD	n	N	Age $\pm 2\sigma$ (Ma)	K/Ca	Age $\pm 2\sigma$ (Ma)	$^{40}\text{Ar}/^{36}\text{Ar}$ Ic	MSWD
FV23	16D08339	gm	334.4 \pm 4.2 Ka	100.00	0.169	1.13	31	31	334.4 \pm 4.2 Ka	0.347	0.34 \pm 0.01	294.85	1.14
FV21	16D08292	gm	2.08 \pm 0.01	61.80	0.102	1.41	15	31	2.08 \pm 0.01	0.122	2.05 \pm 0.07	301.72	1.44
FV10	16D08219	gm	2.82 \pm 0.08	50.45	0.016	1.57	15	31	2.56 \pm 0.06	0.023	2.65 \pm 0.16	298.86	1.17
FV6	16D08172	gm	3.52 \pm 0.04	25.72	0.037	1.01	8	31	3.54 \pm 0.04	0.062	3.80 \pm 0.32	279.36	0.72
FV1	16D08125	gm	3.51 \pm 0.02	97.00	0.030	0.81	27	31	3.65 \pm 0.03	0.033	3.51 \pm 0.04	295.27	0.84
FU18	16D08386	gm	3.35 \pm 0.02	97.47	0.034	0.76	28	31	3.49 \pm 0.03	0.038	3.33 \pm 0.03	297.40	0.67
FT38	16D08078	gm	2.69 \pm 0.05	57.10	0.024	1.02	13	31	2.85 \pm 0.06	0.022	2.60 \pm 0.13	300.05	0.90
FT27	16D08005	gm	3.68 \pm 0.04	92.56	0.022	0.60	24	31	3.95 \pm 0.05	0.019	3.68 \pm 0.05	295.74	0.63
FT8	16D07958	gm	4.26 \pm 0.03	85.19	0.043	0.93	23	31	4.20 \pm 0.04	0.034	4.28 \pm 0.05	294.33	0.94
FT4	16D07911	gm	4.20 \pm 0.02	78.44	0.050	1.19	18	31	4.15 \pm 0.02	0.054	4.22 \pm 0.05	294.11	1.24
FS23	16D06588	gm	4.38 \pm 0.06	58.92	0.046	2.09	19	32	4.44 \pm 0.05	0.048	4.40 \pm 0.30	295.12	2.21
FS15	16D06513	gm	5.07 \pm 0.05	50.97	0.029	1.16	13	32	5.13 \pm 0.05	0.023	5.10 \pm 0.44	294.48	1.26
FS1	16D07070	gm	5.28 \pm 0.05	42.50	0.075	2.52	12	31	5.39 \pm 0.03	0.069	5.41 \pm 0.14	288.72	1.98
FR26	16D07117	gm	5.33 \pm 0.07	24.56	0.025	1.51	12	31	5.72 \pm 0.02	0.097	5.53 \pm 0.14	292.42	0.78
FR15	16D06466	gm	6.25 \pm 0.02	30.58	0.168	2.56	7	31	6.05 \pm 0.01	0.141	6.25 \pm 0.14	297.58	3.08
FR7	16D06415	gm	6.74 \pm 0.04	30.51	0.165	6.26	6	34	6.39 \pm 0.02	0.093	6.78 \pm 0.17	286.67	7.36
FR4	15D35592	gm	7.00 \pm 0.04	17.89	0.246	1.91	5	32	6.28 \pm 0.05	0.198	6.98 \pm 0.15	297.09	2.43

disturbed ^{39}Ar release patterns reported here for FR15, FR7 and FR4. On this note, it is with some hesitance that we dismiss the legacy ages. In summary, our interpretations of the 17 new $^{40}\text{Ar}/^{39}\text{Ar}$ age determinations show that the composite stratigraphic section from base of FR to top FV span the interval of 7.00 ± 0.04 Ma to 0.33 ± 0.00 Ma, thereby extending the previously established absolute age range (Dagley *et al.* 1967; Watkins & Walker 1977; Mussett *et al.* 1980) by one Myr.

The intersection stratigraphic correlations presented in Fig. 2 are based on the combined observations of lithostratigraphy (distinct traceable lava groups and sedimentary deposits), magnetostratigraphy and absolute age determinations, and builds on the work by Watkins & Walker (1977) and Gudmundsson (1978). Hiatuses or periods characterized by no or little volcanic activity are identified throughout the composite stratigraphy with: (i) major hiatuses (>500 kyr duration) between FV22–23, FV11–12, FV6–9/FUB17–20/FT35–FT37, FS19–21 and between FR16–17, and (ii) minor hiatuses ($\leq c.$ 350 kyr) between FT23–24, FT2–3/FS23–24, FR12–13 and between FR4–5. The major hiatus identified in FV, FU and FT consists of two briefer hiatuses with sporadic volcanic products in-between. The intermittent volcanic products are of variable age and magnetic polarity, reflecting a decrease in lateral continuity in volcanic units during the late Pliocene and Pleistocene because of increased topographic relief, (glacial) erosion and interaction between volcanism and ice cover. In the discussion we present a magneto-chronostratigraphic age model for the composite stratigraphic section in Fljótisdalur.

4.2 Palaeodirections

We analysed the demagnetization data using Puffinplot (Lurcock & Wilson 2012) and determined the direction of the characteristic remanent magnetization (ChRM) using principal component analysis (Kirschvink 1980). Five or more points were used and only

directions trending to the origin with a maximum angle of deviation (MAD) $<5^\circ$ were accepted. Two sites were discarded from the analysis due poor demagnetization results. A total of 1165 independent directions were accepted from 123 sites, with mean, minimum (min) and maximum (max) number (N) of independent ChRMs per site $N_{\text{mean}} = 9.5$, $N_{\text{min}} = 5$ and $N_{\text{max}} = 14$, respectively (Fig. 5a). Typical stepwise AF- and thermal-demagnetization plots are shown in Fig. 5(b)–(g). As shown, the magnetic remanence generally resides in a stable single component of magnetization acquired during original cooling. A secondary viscous component of smaller magnitude was generally removed by AF-demagnetization peak fields of 5–10 mT or temperatures of $\sim 250^\circ\text{C}$. These results are similar to the observations obtained by Døssing *et al.* (2016) from Jökuldalur and on extrusive basaltic rocks from other palaeomagnetic studies in Iceland (Kristjánsson 2008; Linder & Leonhardt 2009).

We corrected the ChRM directions for a post-emplacement tectonic westerly (azimuth= 280°) dip of 2.5° – 8.0° . For this we used a handful of *in situ* dip measurements and the eastern Iceland dip-versus-age model of Watkins & Walker (1977) using our updated magnetostratigraphic age model for sections FR to FV (see Section 5).

We also computed the site palaeolatitude by correcting the present-day site latitudes for a (very) small post-emplacement plate kinematic movement. We used the NNR-MORVEL56 plate motion model of Argus *et al.* (2011), which is a set of angular velocities relative to the unique reference frame in which there is no net rotation of the lithosphere. According to NNR-MORVEL56, eastern Iceland (Eurasia Plate) is currently moving in a north-northeasterly direction with a speed and azimuth of 18.28 mm yr^{-1} and 26.98 , respectively. This movement transfers into a northward velocity of 16.29 mm yr^{-1} and an eastward velocity 8.29 mm yr^{-1} as obtained using the online plate motion calculator of UNAVCO (<https://www.unavco.org/software/geodetic-utilities/plate-motion-calculator/plate-motion-calculator.html>). The plate kinematic correction implies that the true site latitude and

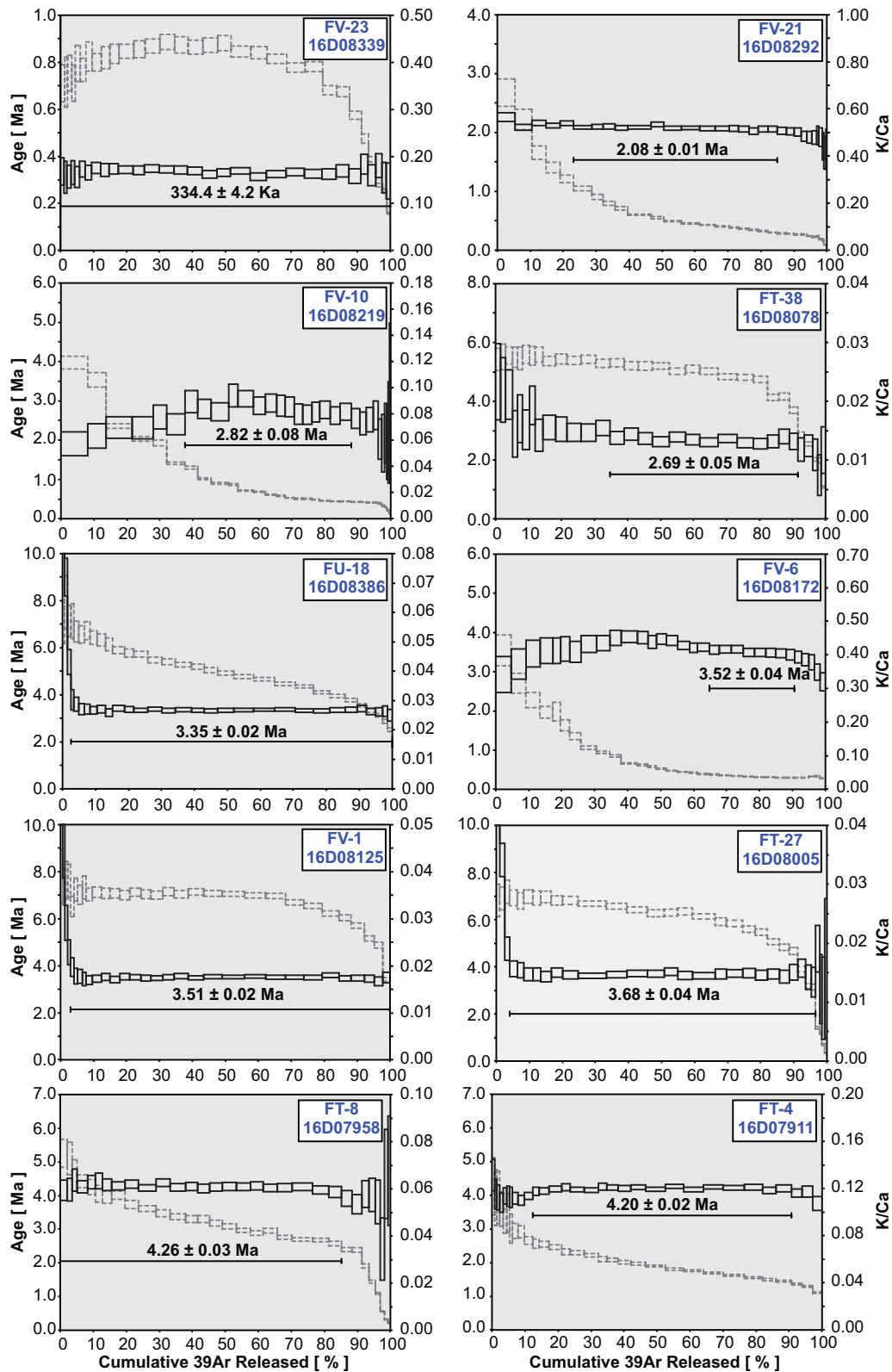


Figure 4. High-resolution incremental heating $^{40}\text{Ar}/^{39}\text{Ar}$ age spectra on groundmass separates subaerial basalt samples from profiles FV, FU, FT, FS and FR in Fljótssdalur, eastern Iceland. The age spectra are presented in stratigraphic order from young to old in the lava pile. The $^{40}\text{Ar}/^{39}\text{Ar}$ ages are weighted age estimates with errors reported at the 95 percent confidence level, including 0.2–0.3 percent standard deviations in the J -value. All samples were monitored against FCT sanidine (28.201 ± 0.023 Ma, 1σ) as calibrated by Kuiper *et al.* (2008). Solid black lines are $^{40}\text{Ar}/^{39}\text{Ar}$ ages and stippled gray lines are K/Ca values. Data are listed in Table 1 and ArArCALC age calculation files will be made available for download from the EarthRef.org Digital Archive (ERDA) and GEOCHRON.

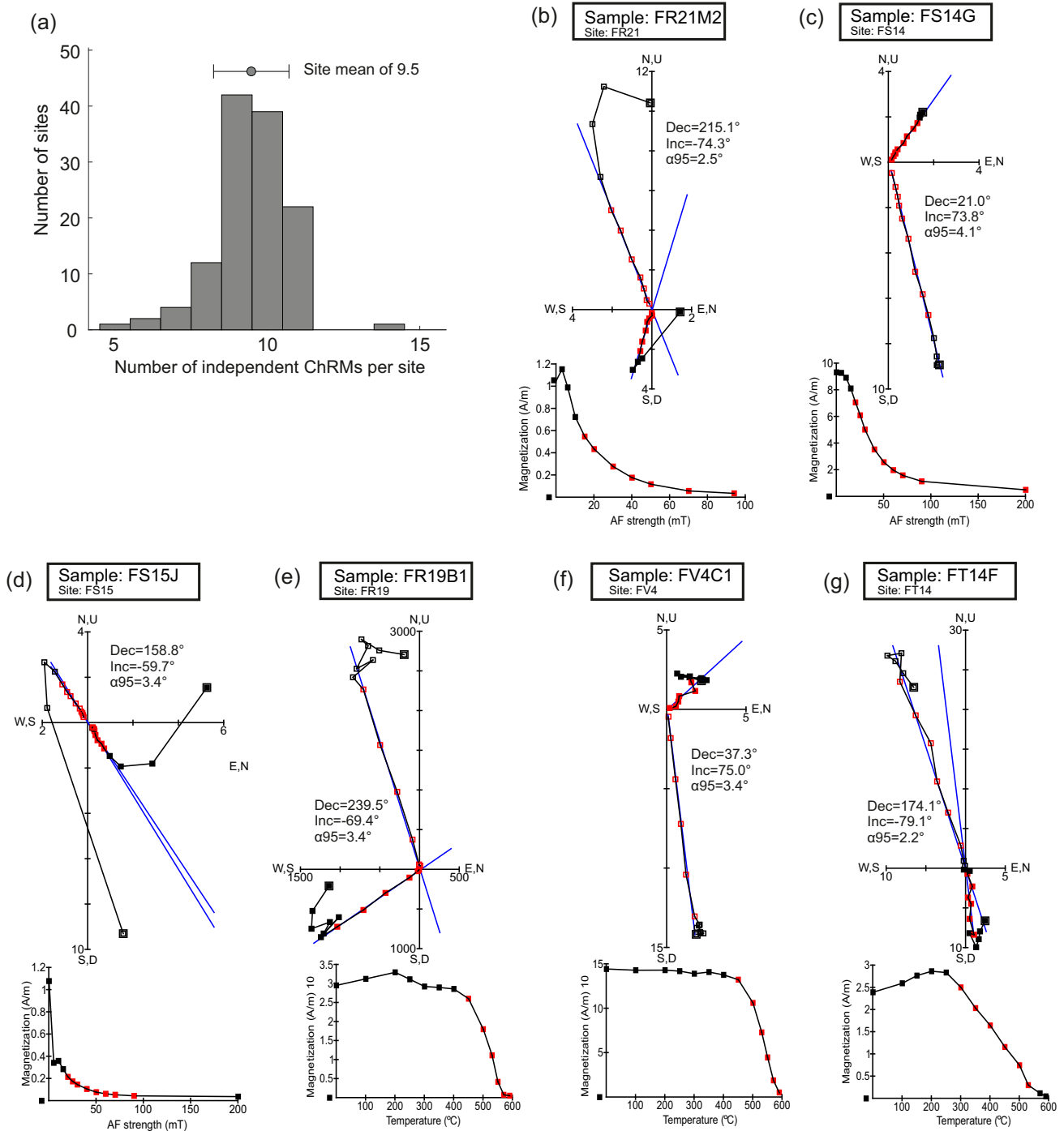


Figure 5. (a) Histogram of independent successful ChRM directions per site for the accepted 123 sites. (b)–(g) Examples of typical evolution of direction and intensity of the natural remanent magnetization during AF demagnetization (b–d) and thermal demagnetization (e–g).

longitude at the time of the rock formation are located to the southwest of the present-day site location. For the oldest site, FR4, the plate kinematic latitude correction is -1.0 , while the longitude correction is -1.2 . For this study, we assume that the plate motions defined by the NNR-MORVEL56 model are valid for the period represented by sites in this study.

The 123 accepted mean site directions were calculated using Fisher (1953) statistics by averaging AF and thermal results. The results are shown in Table 2 together with the corresponding VGPs. A summary of the site mean statistics, including n , k ,

Declination, Inclination, VGP latitude and polarity corrected for post-emplacement tectonic movements, is shown for each section in Fig. 6. In Fig. 7, we plot the site mean directions and VGPs for data corrected for post-emplacement tectonic movements (Supporting Information Fig. S2-3 shows the corresponding plot for data not corrected for post-emplacement tectonic movements).

As shown in Figs 6 and 7 we obtain well-grouped ChRM directions for all sites with $n_{\text{mean}} = 9.5$ ($n_{\text{min}} = 5$, $n_{\text{max}} = 14$), $k_{\text{mean}} = 353$ ($k_{\text{min}} = 47$, $k_{\text{max}} = 1383$) and $\alpha_{95_{\text{mean}}} = 3.1^\circ$ ($\alpha_{95_{\text{min}}} = 1.4^\circ$, $\alpha_{95_{\text{max}}} = 9.6^\circ$), where k is the precision parameter and α_{95}

Table 2. Fljótisdalur palaeodirectional site statistics. Site: Site name. Elev: GPS-altitudes (m). Azi/Dip: Azimuth(°) and dip(°) of bedding. SLon(°N)/SLat(°E): Site Longitude/Latitude (°). SLon(°N)/SLat(°E): Site Longitude/Latitude (°) corrected for plate motions using the NNR-MORVEL56 model (Argus et al. 2011). N: Total number of independent ChRMs per site. D/I: Site mean declination(°)/inclination(°). D'/I': Site mean declination(°)/inclination(°) corrected for bedding tilt. α_{95} (°) Fisher (1953) circle of 95% confidence. *k*: Estimate of the Fisher (1953) precision parameter. *R*: Length of the vector sum, or resultant vector *R* (Fisher 1953). VLon/VLat: Virtual geomagnetic pole position (°). *dp/dm*: 95% confidence angles(°) in parallel and meridian. *Pol*: Polarity. *MAge*: Model age(Ma)/Chron (B, Brunhes; M, Matuyama; Ga, Gauss, Gi, Gilbert, Mi, Miocene)—see Section 5. Sites removed for further analysis due to statistical overlap with sites immediately above/below or comparable sites in other sections: FR16, FR22, FS2, FS29, FT37, FUB17, FU22, FV4, FV6, FV9.

Site	Elev	Azi/Dip	SLon	SLat	SLon'	SLat'	N	D	I	D'	I'	α_{95}	k	R	VLon	VLat	dp/dm	Pol.	MAge
FV23	680	280/0.0	-15.39350	64.88557	-15.45956	64.83145	10	254	88	254	88	3.0	267	10	337	64	5.9/5.9	N	0.37/B
FV22	653	280/2.5	-15.38992	64.88903	-15.75638	64.58416	11	218	-70	212	-71	2.9	252	11	269	-72	4.4/5	R	2.08/M
FV21	644	280/2.6	-15.38600	64.88933	-15.75438	64.58249	11	233	-74	226	-76	2.5	333	11	239	-70	4.2/4.6	R	2.10/M
FV20	628	280/2.7	-15.38025	64.88880	-15.75193	64.57954	10	208	-69	202	-70	2.7	315	10	290	-74	4/4.7	R	2.12/M
FV19	619	280/2.7	-15.37763	64.88877	-15.75193	64.57787	8	218	-78	206	-79	4.8	132	8	217	-79	8.8/9.2	R	2.13/M
FV18	603	280/2.8	-15.37145	64.88752	-15.74848	64.57392	9	194	-74	184	-74	1.9	772	9	320	-85	3/3.3	R	2.15/M
FV17	599	280/2.8	-15.37118	64.88745	-15.74937	64.57218	9	287	-85	296	-88	3.1	274	9	172	-63	6.2/6.2	R	2.15/M
FV16	597	280/2.9	-15.37037	64.88745	-15.74882	64.57181	10	202	-81	184	-81	2.6	345	10	172	-82	4.9/5	R	2.16/M
FV15	590	280/2.9	-15.36902	64.88747	-15.74937	64.57052	10	170	-87	138	-85	2.5	369	10	143	-71	4.9/5	R	2.16/M
FV14	582	280/2.9	-15.36848	64.88752	-15.75015	64.57004	10	135	-74	130	-72	3.2	224	10	80	-65	5/5.7	R	2.17/M
FV13	579	280/3.0	-15.36658	64.88683	-15.74981	64.56848	9	235	-71	229	-73	1.4	1384	9	245	-67	2.2/2.5	R	2.18/M
FV12	574	280/3.0	-15.36557	64.88638	-15.74992	64.56656	11	234	-68	228	-70	2.9	249	11	256	-64	4.3/5	R	2.18/M
FV11	559	280/3.1	-15.36282	64.88582	-15.85091	64.48004	10	193	83	216	82	2.6	349	10	330	51	4.9/5	N	2.78/Ga
FV10	550	280/3.1	-15.35525	64.88577	-15.84491	64.47838	11	55	76	45	78	2.3	411	11	45	72	4/4.3	N	2.79/Ga
FV9	534	280/3.3	-15.34385	64.88967	-15.83613	64.48053	9	24	67	16	68	1.6	1016	9	125	74	2.3/2.7	N	2.80/Ga
FV8	526	280/3.3	-15.33105	64.88898	-15.87949	64.43263	9	41	75	29	77	2.4	482	9	61	78	4/4.4	N	3.12/Ga
FV7	520	280/3.3	-15.32545	64.89097	-15.87838	64.43057	8	13	68	4	68	1.9	821	8	152	76	2.7/3.2	N	3.15/Ga
FV6	447	280/3.8	-15.32338	64.89190	-15.94924	64.37094	8	16	78	358	78	3.9	203	8	329	87	6.9/7.3	N	3.56/Gi
FV5	437	280/3.9	-15.32447	64.89220	-15.95246	64.36909	9	18	77	1	77	2.8	330	9	6	89	4.9/5.3	N	3.58/Gi
FV4	433	280/3.9	-15.32173	64.89175	-15.95135	64.36835	9	37	75	23	76	3.4	234	9	65	80	5.8/6.2	N	3.58/Gi
FV3	429	280/3.9	-15.32192	64.89182	-15.95224	64.36761	9	32	75	17	76	2.6	408	9	77	82	4.3/4.7	N	3.59/Gi
FV2	422	280/4.0	-15.31980	64.89117	-15.95179	64.36532	10	52	87	337	87	2.1	510	10	337	70	4.2/4.3	N	3.59/Gi
FV1	415	280/4.0	-15.31887	64.89122	-15.95235	64.36403	8	340	80	324	77	2.9	368	8	275	75	5/5.4	N	3.60/Gi
FU22	585	280/3.5	-15.28667	64.92367	-15.77135	64.52100	9	38	66	31	67	2.6	408	9	102	68	3.5/4.2	N	2.76/Ga
FU21	562	280/3.6	-15.28242	64.92143	-15.77413	64.51153	10	22	65	14	65	2.7	324	10	133	71	3.5/4.4	N	2.80/Ga
FU20	542	280/3.7	-15.28193	64.92042	-15.77635	64.50868	10	349	79	333	78	3.0	264	10	274	79	5.2/5.6	N	2.81/Ga
FU19	535	280/3.8	-15.28102	64.91925	-15.86872	64.42999	11	128	-53	126	-50	2.5	328	11	56	-43	2.2/3.4	R	3.34/Ga
FU18	530	280/3.8	-15.27855	64.91848	-15.86783	64.42807	9	287	-62	289	-66	3.6	201	9	215	-36	4.9/6	R	3.35/Ga
FUB17	502	280/3.9	-15.27740	64.91600	-15.90324	64.39494	11	348	80	332	78	2.9	249	11	277	78	5.1/5.4	N	3.56/Ga
FU17	487	280/4.0	-15.27733	64.91462	-15.90657	64.39117	10	360	78	343	77	2.5	384	10	263	83	4.3/4.6	N	3.58/Gi
FT38	598	280/3.0	-15.26613	64.94210	-15.75035	64.53900	5	37	68	30	69	3.5	483	5	98	70	5/5.9	N	2.76/Ga
FT37	585	280/3.1	-15.25433	64.94293	-15.74613	64.53353	8	27	64	21	65	2.8	406	8	122	69	3.6/4.4	N	2.80/Ga
FT36	579	280/3.1	-15.24445	64.94300	-15.78627	64.49181	10	156	-78	145	-77	1.9	667	10	90	-75	3.2/3.5	R	3.09/Ga
FT35	512	280/3.6	-15.21183	64.94888	-15.83469	64.43090	9	12	84	342	83	2.6	395	9	325	77	5/5.1	N	3.54/Ga
FT34	496	280/3.7	-15.21082	64.94815	-15.83724	64.42694	10	10	73	358	73	3.2	225	10	173	84	5.1/5.7	N	3.56/Ga
FT33	490	280/3.7	-15.21037	64.94782	-15.83757	64.42583	9	84	80	75	83	2.0	654	9	15	65	3.9/3.9	N	3.57/Gi
FT32	484	280/3.8	-15.20927	64.94767	-15.83791	64.42472	11	324	60	320	57	2.2	446	11	225	55	2.3/3.2	N	3.58/Gi
FT31	479	280/3.8	-15.20887	64.94755	-15.83902	64.42380	9	345	84	322	81	2.6	390	9	303	74	4.9/5	N	3.58/Gi
FT30	460	280/3.9	-15.20768	64.94743	-15.84224	64.41929	11	304	74	300	71	2.1	481	11	265	59	3.1/3.6	N	3.61/Gi
FT29	446	280/4.0	-15.20573	64.94803	-15.84335	64.41770	10	68	69	61	72	4.0	147	10	60	60	6.2/7	N	3.63/Gi
FT28	439	280/4.1	-15.20563	64.94795	-15.84490	64.41641	10	234	75	243	72	2.8	308	10	305	41	4.3/4.9	N	3.63/Gi
FT27	420	280/4.2	-15.20387	64.94783	-15.84712	64.41289	11	235	-71	225	-74	3.3	198	11	244	-69	5.3/5.9	R	3.66/Gi
FT26	416	280/4.2	-15.20415	64.94813	-15.84801	64.41215	8	239	-68	232	-71	3.8	209	8	250	-63	5.8/6.7	R	3.66/Gi
FT25	398	280/4.4	-15.20352	64.94825	-15.85201	64.40883	9	198	-82	170	-81	2.7	371	9	144	-81	5/5.2	R	3.69/Gi
FT24	380	280/4.5	-15.20222	64.94828	-15.85401	64.40550	9	170	-46	166	-44	2.6	399	9	5	-51	2/3.2	R	3.71/Gi
FT23	343	280/4.7	-15.19855	64.94915	-15.91371	64.35433	9	169	-62	161	-60	1.9	750	9	18	-65	2.2/2.9	R	4.07/Gi
FT22	337	280/4.8	-15.19777	64.94907	-15.91405	64.35352	9	122	-76	116	-72	2.9	317	9	90	-59	4.5/5.1	R	4.07/Gi
FT21	325	280/4.8	-15.19763	64.94895	-15.91671	64.35100	10	203	-80	174	-80	2.8	300	10	149	-83	5.1/5.4	R	4.09/Gi
FT20	322	280/4.9	-15.19740	64.94880	-15.91638	64.35045	9	229	-70	217	-73	3.4	233	9	256	-72	5.4/6	R	4.09/Gi
FT19	309	280/5.0	-15.19758	64.94870	-15.91871	64.34934	9	245	-67	237	-71	3.9	176	9	246	-61	5.9/6.8	R	4.10/Gi
FT18	314	280/4.9	-15.19678	64.94875	-15.91860	64.34860	10	182	-82	153	-80	3.2	233	10	116	-78	5.8/6.1	R	4.11/Gi
FT17	298	280/5.0	-15.19682	64.94848	-15.92015	64.34630	11	296	-88	87	-87	3.6	160	11	151	-64	7.2/7.2	R	4.11/Gi
FT16	291	280/5.1	-15.19673	64.94847	-15.92126	64.34538	9	192	-66	181	-66	3.8	189	9	343	-74	5/6.1	R	4.12/Gi
FT15	312	280/4.9	-15.19657	64.94840	-15.92237	64.34446	10	170	-81	147	-78	2.5	387	10	102	-76	4.4/4.6	R	4.13/Gi
FT14	289	280/5.1	-15.19607	64.94833	-15.92471	64.34168	9	174	-79	153	-77	2.2	554	9	89	-78	3.8/4.1	R	4.15/Gi
FT13	268	280/5.2	-15.19570	64.94793	-15.92737	64.33947	11	319	-82	1	-85	2.7	297	11	164	-54	5.2/5.2	R	4.16/Gi
FT12	260	280/5.3	-15.19543	64.94790	-15.92815	64.33799	9	239	-67	229	-71	3.9	175	9	252	-65	5.9/6.8	R	4.17/Gi
FT11	253	280/5.3	-15.19525	64.94785	-15.92970	64.33670	7	224	-77	200	-79	5.4	127	7	212	-81	9.8/10.2	R	4.18/Gi
FT10	245	280/5.4	-15.19562	64.94785	-15.93248	64.33522	9	230	-73	213	-76	4.7	123	9	242	-76	7.9/8.6	R	4.19/Gi
FT9	237	280/5.4	-15.19522	64.94750	-15.93326	64.33374	10	139	-62	134	-57	2.4	416	10	53	-52	2.5/3.5	R	4.20/Gi
FT8	226	280/5.5	-15.19435	64.94765	-15.93470	64.33171	10	355	67	343	65	2.9	287	10	201	71	3.8/4.6	N	4.21/Gi
FTB7	222	280/5.5	-15.19442	64.94760	-15.93559	64.33097	10	9	86	318	83	2.7	315	10	312	72	5.2/5.3	N	4.22/Gi
FT6	206	280/5.7	-15.19355	64.94728	-15.94092	64.32553	8	40	66	28	68	3.5	246	8	102	71	5/6	N	4.25/Gi
FT4	170	280/5.9	-15.19263	64.94700	-15.9														

Table 2. Continued

Site	Elev	Azi/Dip	SLon	SLat	SLon'	SLat'	N	D	I	D'	I'	$\alpha 95$	k	R	VLon	VLat	dp/dm	Pol.	MAge
FS24	364	280/5.5	-15.11562	64.98293	-15.86936	64.35617	11	66	63	59	68	3.0	238	11	70	57	4.1/5	N	4.29/Gi
FS23	350	280/5.6	-15.11527	64.98245	-15.88511	64.34123	10	220	-54	212	-57	2.1	555	10	293	-57	2.2/3	R	4.38/Gi
FS22	330	280/5.7	-15.11465	64.98237	-15.88756	64.33920	11	245	-57	239	-61	1.7	736	11	259	-50	2.2/6	R	4.40/Gi
FS21	320	280/5.8	-15.11430	64.98223	-15.89011	64.33624	10	223	-68	209	-70	3.9	152	10	275	-72	5.8/6.8	R	4.42/Gi
FS20	303	280/5.8	-15.11402	64.98193	-15.96813	64.27133	10	340	77	323	73	2.6	342	10	255	72	4.2/4.7	N	4.86/Gi
FS19	300	280/5.9	-15.11402	64.98192	-15.99560	64.24847	10	218	-82	176	-82	2.6	344	10	158	-79	5/5.1	R	5.02/Gi
FS18	294	280/5.9	-15.11358	64.98177	-15.99760	64.24680	9	307	-65	315	-70	2.7	372	9	194	-33	4/4.6	R	5.03/Gi
FS17	285	280/5.9	-15.11325	64.98162	-15.99860	64.24514	10	177	-67	164	-65	3.3	210	10	19	-71	4.4/5.4	R	5.04/Gi
FS16	276	280/6.0	-15.11323	64.98148	-15.99971	64.24322	11	161	-67	150	-63	1.8	615	11	39	-65	2.3/2.9	R	5.04/Gi
FS15	271	280/6.0	-15.11318	64.98137	-16.00393	64.23971	9	159	-60	151	-56	3.4	225	9	30	-58	3.6/5	R	5.07/Gi
FS14	252	280/6.1	-15.11277	64.98103	-16.00637	64.23767	10	21	74	360	74	4.1	140	10	166	86	6.6/7.4	N	5.08/Gi
FS13	241	280/6.2	-15.11258	64.98083	-16.00815	64.23619	9	43	65	30	68	3.0	301	9	99	69	4.2/5	N	5.09/Gi
FS12	233	280/6.2	-15.11237	64.98065	-16.01270	64.23157	10	166	82	210	82	3.5	192	10	332	50	6.6/6.8	N	5.12/Gi
FS11	208	280/6.4	-15.11157	64.98028	-16.01604	64.22780	9	259	53	261	47	3.5	224	9	275	22	2.9/4.5	N	5.14/Gi
FS10	193	280/6.4	-15.11122	64.98003	-16.01681	64.22632	10	29	57	19	58	2.0	587	10	131	62	2.2/3	N	5.15/Gi
FS9	185	280/6.5	-15.11115	64.97990	-16.02103	64.22281	9	26	68	10	69	4.1	159	9	136	78	6/7	N	5.18/Gi
FS8	166	280/6.6	-15.11098	64.97965	-16.02281	64.22133	9	358	86	306	82	2.1	585	9	305	69	4/4.1	N	5.19/Gi
FS7	157	280/6.6	-15.11043	64.97950	-16.02248	64.22078	9	25	88	300	84	2.5	427	9	314	68	4.8/4.9	N	5.19/Gi
FS6	156	280/6.6	-15.11067	64.97940	-16.02525	64.21830	9	344	86	305	81	2.7	367	9	298	69	5/5.2	N	5.20/Gi
FS5	147	280/6.7	-15.11072	64.97907	-16.02992	64.21442	11	11	68	355	67	2.8	265	11	177	75	3.9/4.7	N	5.23/Gi
FS4	124	280/6.8	-15.10990	64.97848	-16.02980	64.21268	9	90	80	73	86	2.6	382	9	2	65	5.2/5.2	N	5.23/Gi
FS3	123	280/6.8	-15.10968	64.97827	-16.03314	64.20991	10	34	70	15	72	1.8	710	10	111	80	2.8/3.2	N	5.25/Gi
FS2	107	280/6.9	-15.10910	64.97797	-16.03680	64.20603	9	62	71	45	76	2.1	592	9	56	71	3.6/3.9	N	5.28/Gi
FS1	86	280/7.0	-15.10843	64.97767	-16.03802	64.20418	9	356	75	335	72	3.5	223	9	234	76	5.3/6.1	N	5.29/Gi
FR28	411	280/6.0	-15.01633	65.01493	-15.94380	64.24303	11	49	71	32	74	4.4	107	11	73	75	7.3/8	N	5.28/Gi
FR27	393	280/6.1	-15.01642	65.01475	-15.94780	64.23970	11	1	69	347	67	2.9	253	11	197	74	4/4.8	N	5.30/Gi
FR26	378	280/6.2	-15.01565	65.01463	-15.95113	64.23693	9	337	-88	81	-85	3.1	279	9	142	-61	6/6.1	R	5.32/Gi
FR25	364	280/6.3	-15.01562	65.01423	-15.95424	64.23334	10	154	-85	123	-80	3.6	179	10	117	-68	6.7/7	R	5.34/Mi
FR24	353	280/6.3	-15.01570	65.01387	-15.95668	64.23131	10	237	-78	208	-81	2.8	300	10	201	-77	5.2/5.4	R	5.35/Mi
FR23	335	280/6.4	-15.01507	65.01383	-15.95901	64.22854	14	345	-58	355	-61	5.1	62	14	168	-16	5.9/7.8	R	5.37/Mi
FR22	338	280/6.4	-15.01562	65.01328	-15.96090	64.22680	10	230	-76	204	-79	2.8	307	10	219	-79	5/5.2	R	5.38/Mi
FR21	315	280/6.6	-15.01535	65.01332	-15.96412	64.22329	11	215	-74	191	-76	2.6	311	11	262	-85	4.4/4.8	R	5.40/Mi
FR20	308	280/6.6	-15.01495	65.01303	-15.96568	64.22199	10	260	-72	249	-78	2.4	425	10	216	-63	4.2/4.4	R	5.41/Mi
FR19	304	280/6.6	-15.01348	65.01337	-15.96457	64.22126	11	240	-69	225	-74	3.4	184	11	245	-69	5.5/6.1	R	5.41/Mi
FR18	286	280/6.7	-15.01428	65.01285	-15.96956	64.21793	7	237	-75	212	-79	2.6	545	7	218	-77	4.7/4.9	R	5.44/Mi
FR17	258	280/6.9	-15.01415	65.01265	-15.97578	64.21275	11	244	-55	237	-60	1.9	557	11	261	-50	2.2/2.9	R	5.47/Mi
FR16	255	280/6.9	-15.01162	65.01287	-16.10920	64.10008	10	29	-81	59	-77	2.3	443	10	133	-46	4/4.3	R	6.24/Mi
FR15	235	280/7.0	-15.01177	65.01267	-16.11142	64.09823	9	24	-83	63	-79	2.8	346	9	134	-50	5/5.3	R	6.25/Mi
FR14	222	280/7.1	-15.01103	65.01263	-16.11331	64.09583	9	239	-73	217	-78	3.3	240	9	229	-75	5.9/6.2	R	6.27/Mi
FR13	193	280/7.3	-15.01037	65.01250	-16.11675	64.09213	10	229	-69	210	-73	3.2	227	10	261	-75	5.1/5.7	R	6.30/Mi
FR12	186	280/7.3	-15.00977	65.01173	-16.18158	64.03719	9	16	66	360	66	3.1	279	9	165	74	4.1/5	N	6.67/Mi
FR11	181	280/7.3	-15.01005	65.01152	-16.18270	64.03627	11	236	49	241	43	5.5	70	11	290	11	4.2/6.8	N	6.67/Mi
FR10	171	280/7.4	-15.01017	65.01087	-16.18492	64.03342	8	252	72	260	65	6.1	229	4	286	37	8/9.8	N	6.68/Mi
FR9	144	280/7.6	-15.01000	65.01058	-16.19091	64.02843	7	1	83	318	79	8.9	47	7	289	73	16.1/16.9	N	6.72/Mi
FR8	144	280/7.6	-15.01010	65.01037	-16.19136	64.02706	9	37	80	354	80	3.5	216	9	328	83	6.5/6.7	N	6.72/Mi
FR7	130	280/7.6	-15.00968	65.01008	-16.19402	64.02484	8	18	71	356	71	3.1	330	8	180	81	4.6/5.3	N	6.74/Mi
FR6	99	280/7.8	-15.00962	65.00942	-16.20091	64.01811	6	38	77	2	79	3.2	442	6	352	86	5.7/6	N	6.78/Mi
FR5	97	280/7.8	-15.00683	65.01015	-16.19880	64.01837	7	63	73	40	79	3.0	396	7	43	74	5.4/5.7	N	6.78/Mi
FR4	68	280/8.0	-15.00688	65.00922	-16.23819	63.98460	8	218	-65	200	-68	4.9	129	8	295	-73	6.9/8.2	R	7.00/Mi

is the circle of 95 per cent confidence (Fisher 1953). Of the 123 sites, 62 sites are of normal polarity and 61 sites are of reverse polarity. Based on common mean tests (Tauxe 1998), direct field observations and the new radiometric ages (Table 1), we found that only 113 sites are statistically distinct (of the sites that apparently recorded the same snapshot of the field, we simply discarded one of those sites for the further analysis; see Table 2 for information about sites excluded for further analysis). Hence, there exists some stratigraphic overlap between the investigated sections; see Fig. 2. In addition, we mapped 12 potential transitional polarity sites with $-45^\circ < \text{VGP}_{\text{latitude}} < 45^\circ$. Seven of these sites have normal polarity and five have reverse polarity (Table 2).

Sections FR to FV were drilled by Watkins & Walker (1977) who sampled two cores per site. A review of the polarity mapping of Watkins & Walker (1977) shows that there's a high correlation with the polarity distributions found in this study, with only a few major exceptions. Watkins & Walker (1977) suggest, for example, the presence of a reverse polarity near site FV7 and FV8, while this is not observed in this study (Table 2). In section FT, Watkins & Walker (1977) mapped to FT32/FT33 (positive polarity), while additional flows were mapped in this study and revealed the evidence of a short reverse polarity interval at FT35, followed by normal polarities. When looking into detail, however, a marked difference is found

between mapped declination and inclination values for positively correlated flows (see Supporting Information Fig. S2-1). Notably, the differences show a Gaussian distribution, indicating that they relate to higher levels of random noise in the data set of Watkins & Walker (1977) as compared to the new data set. Thus, some of the observed differences may relate to sun versus magnetic bearings (96 per cent of all cores in the new data set were oriented using a sun compass), while other differences may relate to the fact that Watkins & Walker (1977) collected only two cores per site.

For the following analysis we excluded the 12 potential transitional sites as defined by a VGP latitude cut-off of $\pm 45^\circ$ (see Section 5). We further excluded two sites with $k = 46.9$ and $k = 49.9$. This cut-off is based on results by Supakulopas (2018) on Icelandic lavas, showing that the estimated VGP dispersion of a suite becomes stable around a certain mean for $k > \sim 80$, with a marked drop in change between $k > \sim 60$ and $k > \sim 80$. The resulting 46 normal and 53 reverse site mean directions are defined by $N \geq 5$ and $\alpha 95 \leq 6$. As shown in Supporting Information Fig. S2-2 shown to be Fisherian, the hypothesis that the distribution of directions is Fisherian cannot be rejected at a 5 per cent significance level. The 46 normal sites yield a mean direction of declination (D_{normal}) = 0.1° and inclination (I_{normal}) = 77.5° ($\alpha 95 = 3.0^\circ$), while the 53 reverse sites yield a mean direction of declination (D_{reverse}) = 190.4° and

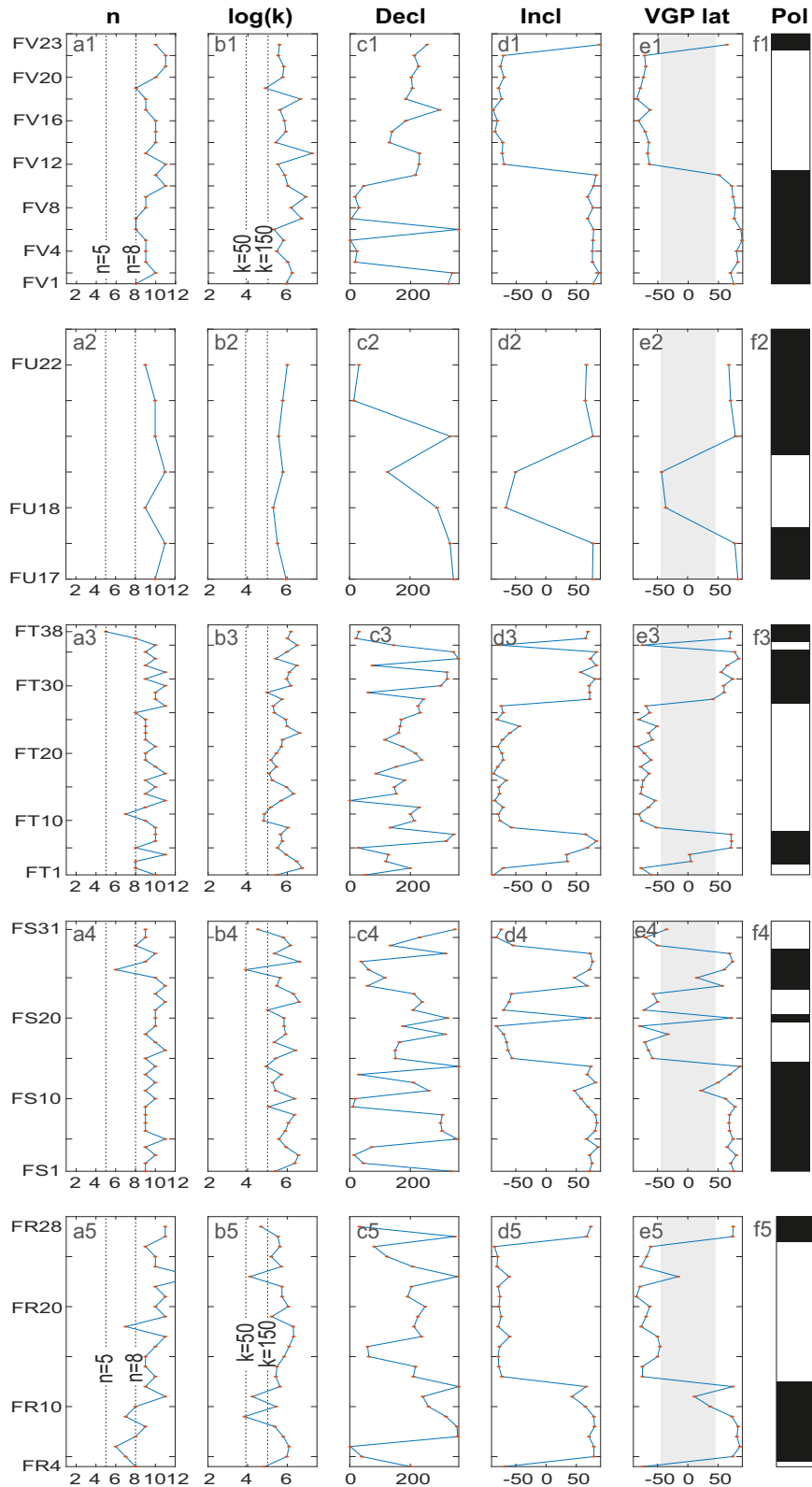


Figure 6. Summary plots of information provided in Table 2 for sections FR (top panels) to FV (bottom panels). Columns: (a) Number of independent directions, ChRM (n); (b) log of Fisher precision parameter k ; (c) declination ($^{\circ}$); (d) Inclination ($^{\circ}$); (e) VGP latitude ($^{\circ}$); (f) polarity. For reference, the vertical lines in Column (a) are for $n = 4$ and $n = 8$, while the vertical lines in (b) are for $k = 50$ and $k = 150$. The grey bar in Column (e) highlights VGP latitudes between -45° and 45° .

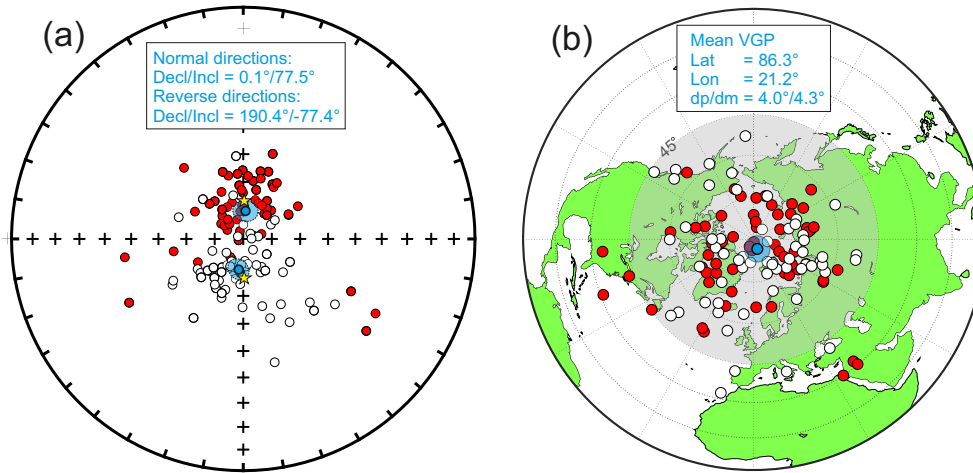


Figure 7. (a) Equal area projection of Fljótisdalur site mean palaeodirections corrected for post-emplacment tectonic movements. Filled and open circles plot on the lower and upper hemisphere, respectively. Grand mean directions (filled blue circles) with $\alpha 95$ confidence ellipses (transparent blue) calculated for sites that fulfill quality criteria (see Table 3). Yellow stars: Expected directions from a GAD field. (b) VGP positions. Transparent grey circle is at 45° latitude. Filled (open) circles are northern hemisphere (antipodes of southern hemisphere) VGP positions. Filled blue circles: Grand mean VGP (palaeomagnetic pole) and 95 per cent confidence ellipses calculated for sites that fulfill quality criteria (see Table 3). As shown in (b) the resulting mean palaeomagnetic pole coincides with the North Pole within the 95 per cent confidence limits.

Table 3. Palaeodirectional mean statistics for selected 99 sites that fulfill the quality criteria (see the text). N , total number independent units; D/I , Grand mean declination($^\circ$)/inclination($^\circ$); $\alpha 95$ ($^\circ$), Fisher (1953) circle of 95% confidence; k , Estimate of the Fisher (1953) precision parameter. VGP_{lat}/VGP_{lon} : Grand mean VGP latitude($^\circ$) and longitude($^\circ$). dp/dm : 95% confidence angles($^\circ$) in parallel and meridian.

	N	D	I	$\alpha 95$	k	VGP_{lat}	VGP_{lon}	dp/dm
Normal	46	0.1	77.5	3.0	50	87.3	-17.1	5.3/5.6
Reverse	53	190.4	-77.4	3.3	34	-84.7	37.5	5.8/6.2
Combined	99	5.6	77.5	2.3	41	86.3	21.2	4.0/4.3

inclination ($I_{reverse}$) = -77.4° ($\alpha 95 = 3.3^\circ$) (Table 3, Fig. 7b). The normal and reverse directions pass a bootstrap reversals test (Tauxe 2010) at the 95 per cent level of confidence (Supporting Information Fig. S2-4) and a Watson (1983) V-test (Supporting Information Fig. S2-5) support they have a common mean direction, indicating that the normal and reverse data can be combined. The 53 reverse directions were therefore inverted and an overall mean was calculated of the combined data set of $N = 99$ with a result ($D = 5.6^\circ$, $I = 77.5^\circ$, $\alpha 95 = 2.3^\circ$) that is not significantly different from that expected from a geocentric axial dipole field ($I_{GAD} = 76.5^\circ$). The resulting palaeomagnetic pole position that would give rise to the site mean magnetic field direction under the assumption of a GAD field is shown in Fig. 7(d) and given by $VGP_{latitude}$ and $VGP_{longitude}$ of $86.3^\circ N$ and $21.2^\circ E$ with semi-axes of the ellipse of confidence (dp/dm) of $4.0/4.3$. This result from the Fljótisdalur samples show that the North Pole is within the 95 per cent confidence region of the palaeomagnetic pole.

5 DISCUSSION

5.1 Chronostratigraphic age model of Fljótisdalur

In order to analyse the PSV for specific time intervals, we need a solid age model for sections FR to FV. We have built a new composite chronostratigraphic age model based on field observations (Fig. 2), $^{40}\text{Ar}/^{39}\text{Ar}$ results (Table 1), palaeodirectional data (Table 2) and resultant intersection stratigraphic correlations. The age

model provides a time constraint on our subsequent PSV analysis of the palaeodirectional data.

In Fig. 8(a), we plot the 17 new $^{40}\text{Ar}/^{39}\text{Ar}$ radiometric ages against the cumulative stratigraphic height (CSH) for the dated lava flows in sections FR to FV. We estimated the CSH using the approach of Døssing *et al.* (2016) from Jökuldalur, that is, by considering the stratigraphic thickness of each lava flow and taking into account the tectonic tilt (Table 2) and the relative age chronology of the sections (Fig. 2). Fig. 8(a) indicates that the 17 new $^{40}\text{Ar}/^{39}\text{Ar}$ ages define an overall linear trend with CSH, thus supporting the conclusion of Watkins & Walker (1977); McDougall *et al.* (1977) that the volcanic build-up rate in eastern Iceland can be assumed to have remained overall constant for interhiatus intervals on timescales of 10^5 – 10^6 years. We therefore assumed a segmented linear regression age model as a valid first order approximation to model the $^{40}\text{Ar}/^{39}\text{Ar}$ ages versus CSH. First, we calculated a linear regression model for the long age interval defined by sites FR17 (CSH ~ 190 m) to site FS19 (CSH ~ 570 m) (Fig. 2, Table 2). This interval contains three new $^{40}\text{Ar}/^{39}\text{Ar}$ ages and is bounded above and below by major hiatuses. We used the FR17–FS19 linear regression model as an approximation to the remaining age intervals using the assumption of overall constant build-up rates between major hiatuses. However, for the relatively young and short interval defined by sites FV12–FV22, which was formed during the first glaciations and contains several hiati (Fig. 2), we calculated a separate—and less steep—regression model to fit the $^{40}\text{Ar}/^{39}\text{Ar}$ age of FV21 and the GTS2012 polarity timescale of Gradstein *et al.* (2012) (see below). The result is shown in Fig. 8(b). In order to assign ages to all sites along the sections, we used the age-versus-CSH regression model together with the CSH for all sites to calculate the interpolated model age for each site (Fig. 8c). Finally, we used these model ages to correlate and fit with the polarity intervals of GTS2012 (Gradstein *et al.* 2012) in order to obtain a complete chronostratigraphic age model for Fljótisdalur (Fig. 8d). All model ages are listed in Table 2.

The new age model for sections FR to FV in Fljótisdalur shows that the sections (excluding site FV23 in top of section FV) span the age interval 2.08–7.0 Ma, covering eight normal polarity intervals and ten reverse polarity intervals. The uppermost site, FV23, has a

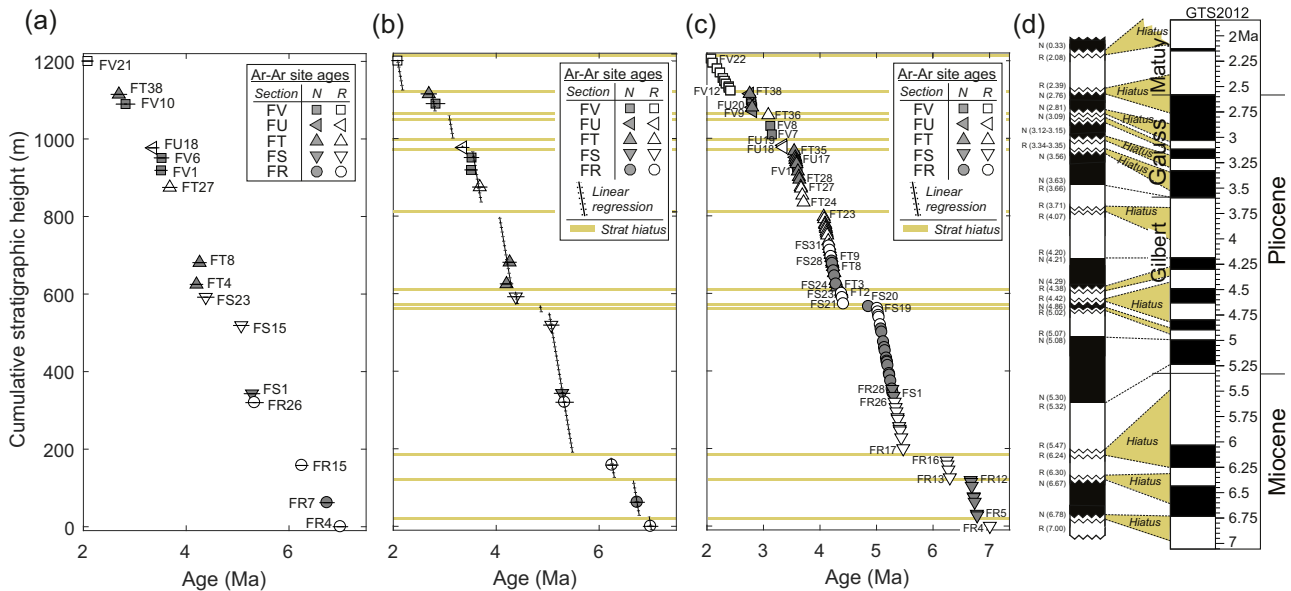


Figure 8. (a) $^{40}\text{Ar}/^{39}\text{Ar}$ ages versus cumulative stratigraphic height (CSH) for sections FR–FV. (b) First-order segmented regression model based on the $^{40}\text{Ar}/^{39}\text{Ar}$ data. Also shown are major hiatuses in the stratigraphic column as determined from field observations (Fig. 2) and $^{40}\text{Ar}/^{39}\text{Ar}$ data (Table 1). (c) Interpolated model ages for all sites according to their CSH. Black (white) symbols are normal (reverse) polarity. (d) Revised magneto-chronostratigraphic age model based on correlation to GTS2012 (Gradstein *et al.* 2012).

radiometric age of 0.33 ± 0.01 Ma and a normal polarity (Tables 1 and 2), indicating that it belongs to Brunhes. Hence, a significant hiatus of 1.75 Myr (0.33–2.08 Ma) exists in the top of Section FV. Additional significant hiatuses in the stratigraphic record are found at the base of site FV12 (2.18–2.78 Ma), at the base of site FS21 (4.41–4.86 Ma) and at the base of site FR17 (5.47–6.24 Ma) (Fig. 2). Although numerous smaller hiatuses are interpreted throughout the sections, the new age model for Fljótisdalur indicate that the 113 distinct flows are well distributed over an age interval of 2.08 Ma to 7.00 Ma (Fig. 8d).

5.2 PSV analysis: 1–7 Ma

A key objective of GAD studies is to estimate TAF and PSV for well-dated time intervals using modern-standard high-quality data (Johnson *et al.* 2008; Cromwell *et al.* 2018). It is also of particular interest to compare normal and reverse polarity intervals, such as the Brunhes and Matuyama chrons, which have repeatedly been shown to record significantly different overall geomagnetic field behaviours (Merrill & McElhinny 1977; Johnson *et al.* 2008; Ziegler *et al.* 2011; Vişan *et al.* 2016; Cromwell *et al.* 2018). However, global palaeomagnetic databases are dominated by results from Brunhes as well as low- to mid-latitudes, whereas data from the older chrons, including the Matuyama, Gauss and Gilbert chrons, as well as from high latitudes, are relatively scarce (Johnson *et al.* 2008; Cromwell *et al.* 2018). Only Lawrence *et al.* (2009) published a large data set for high southerly latitudes that spans the last 7 Myr.

Døssing *et al.* (2016) published modern-standard ($N_{\text{mean}} > 10$ and $k_{\text{mean}} > 200$) palaeodirectional results from 45 individual lava flows with well-constrained $^{40}\text{Ar}/^{39}\text{Ar}$ ages from Jökuldalur (see Fig. 1 for location). Of the 45 lava flows from Jökuldalur, only a single flow is of Brunhes age (0.63 ± 0.03 Ma) while the remaining 44 flows span an age interval of 1.01–3.13 Ma. Eleven of these flows are older than 2 Ma and, hence, have a temporal overlap with the >2.08 Ma flows in sections FR to FV in Fljótisdalur. The temporal spread of the combined Jökuldalur and Fljótisdalur palaeodirectional data

is plotted in Fig. 9. As shown, the Jökuldalur data cluster mainly in the Matuyama reverse polarity chron (1.00–2.59 Ma), whereas the Fljótisdalur data cluster in the Gauss normal and Gilbert reverse polarity chrons (2.59–5.32 Ma) as well as in the late Miocene. Looking at one Myr intervals, the combined data set comprise two directions from 0–1 Ma, 31 directions from 1–2 Ma, 28 directions from 2–3 Ma, 19 directions from 3–4 Ma, 28 directions from 4–5 Ma, 26 directions from 5–6 Ma, and 9 directions from 6–7 Ma, that is, only the 0–1 Ma and 6–7 Ma intervals are particularly poorly constrained by data.

For the following analysis of PSV, we combine the palaeodirectional results from Fljótisdalur and Jökuldalur. We exclude the two Brunhes data points as well as two data points with $k \leq 50$. The site statistics for the combined palaeodirections (Supporting Information Fig. S2-6) show that the palaeodirections provide a high-quality comprehensive data set that is suitable for PSV analysis for the 1–7 Ma interval for high northern latitudes.

The summary statistic used to characterize PSV is the dispersion of the site mean directions through time, which can be evaluated quantitatively using VGPs from each distinct site. We use the modified angular standard deviation for the dispersion S_B (Johnson *et al.* 2008), the root-mean-square angular deviation of VGPs about the geographic axis:

$$S_B = \sqrt{\frac{1}{N-1} \sum_{i=1}^N \left(\Delta_i^2 - \frac{S_{w_i}^2}{N_{s_i}} \right)} \quad (1)$$

Δ_i represents the angular deviation of the pole for the i th site from the geographic north pole, N is the number of sites, and S_B represents the geomagnetic signal remaining after correcting for the within-site dispersion S_{w_i} , determined from N_{s_i} samples.

We calculate and compare the dispersion S_B specifically for the Matuyama ($S_{B(\text{Mat})}$), Gauss ($S_{B(\text{Gauss})}$) and Gilbert ($S_{B(\text{Gil})}$) chrons in order to map any differences in PSV between these geomagnetic intervals such as has been suggested for the Brunhes normal and Matuyama reverse intervals (Johnson *et al.* 2008; Vişan *et al.* 2016;

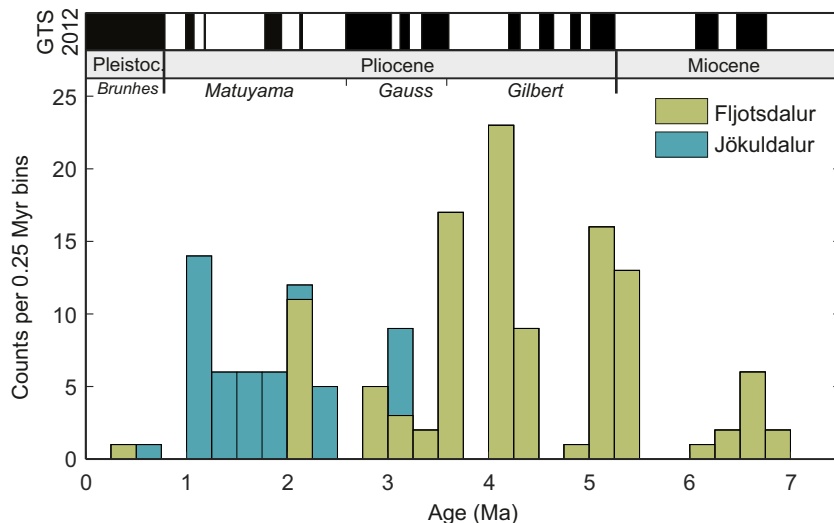


Figure 9. Histogram plot of number of independent site mean directions as a function of model ages for Fljotsdalur data (this study) and Jökuldalur data (Døssing *et al.* 2016). The data have been binned in 0.25 Myr intervals.

Cromwell *et al.* 2018). We further calculate the dispersion for all data combined ($S_{B(1-7)}$) as well as separately for all normal polarity data only ($S_{B(1-7(N))}$) and for all reverse polarity data only ($S_{B(1-7(R))}$). Transitional data are often defined as measurements with an absolute value of VGP latitude less than 45° (Johnson *et al.* 2008) or less than an empirically defined latitude (Vandamme 1994). Based on these latitude cut-off criteria, VGP dispersion values are typically presented for all data (i.e. including low latitude VGPs) as well as for ‘non-transitional’ VGPs only (Johnson *et al.* 2008; Lawrence *et al.* 2009). However, as argued by Kristjansson (2013) VGPs from Iceland show a smooth and continuous frequency distribution in latitude, suggesting that secular variation, excursions and transitions could be treated jointly in VGP statistical studies. This observation is partly supported by the distribution of VGPs in this study, which shows that less than 7 percent of VGP latitudes fall below 45° and that the distribution is relatively smooth (see Supporting Information Fig. S2-7). However, for comparison purposes, we present dispersion results without latitude cut-off as well as for $\pm 45^\circ$ and Vandamme (1994) cut-offs. The calculated VGP dispersion results are listed in Table 4.

If no VGP cut-off is applied to the entire 1–7 Ma data set, we observe a higher dispersion for normal polarity VGPs ($30.3^{37.1}_{23.4}$) as compared to the reverse polarity VGPs ($25.0^{28.1}_{22.0}$) although the two estimates of VGP dispersion are not statistically distinguishable. This result potentially contradicts other studies of a somewhat higher VGP dispersion during reverse polarity intervals (Merrill & McElhinny 1977; Johnson *et al.* 2008). However, it may simply relate to the fact that very-low-latitude VGPs in Fljotsdalur are predominantly observed for normal polarity data and, hence, the data play a (too) significant role when estimating the dispersion without cut-off. If applying a VGP cut-off, we find for the 1–7 Ma data set that the dispersion for normal polarity VGPs becomes $19.7^{22.1}_{17.4}$ and $19.7^{21.9}_{17.3}$ for a $\pm 45^\circ$ and a Vandamme (1994) cut-off, respectively. For the reverse data, the corresponding values become $22.0^{23.8}_{19.9}$ and $22.0^{24.0}_{20.0}$, that is, although the upper and lower bounds of the 1–7 Ma VGP dispersion for the normal and reverse VGPs overlap, these results tend to support previous conclusions of a somewhat higher VGP dispersion during reverse polarity intervals (Johnson *et al.* 2008; Cromwell *et al.* 2018).

When comparing individual chrons, we find that the VGP dispersion for the Matuyama reverse chron ($S_{B(\text{Mat})} = 19.8^{21.7}_{17.8}$) is solely directly comparable to the dispersion for the Gauss normal chron ($S_{B(\text{Gauss})} = 20.4^{24.2}_{15.7}$) if applying the $\pm 45^\circ$ cut-off, whereas markedly higher dispersion values are found for the Gauss chron regardless of using no cut-off or the Vandamme (1994) cut-off. It is evident, however, that the high dispersion for the Gauss data could be related to the limited number of data points (20–23) as compared to the Matuyama (54–55). However, recalculating the mean VGP dispersion for 1000 randomly selected subsets of 23 data points of the Matuyama data results in $S_{B(\text{Mat } 23)} = 19.7^{22.7}_{16.6}$, $S_{B(\text{Mat } 23)} = 19.7^{22.8}_{16.6}$ and $S_{B(\text{Mat } 23)} = 19.3^{21.9}_{16.5}$, respectively, for no cut-off, a $\pm 45^\circ$ cut-off and a Vandamme (1994) latitude cut-off of $\pm 49.5^\circ$. Thus, it appears that a higher dispersion—according to our data—may have characterized the Gauss chron. Interestingly, we find that the mean VGP dispersion for the Matuyama data (regardless of VGP cut-off) is smaller than the mean VGP dispersion found in Brunhes-age rocks from Jan Mayen at 71°N north of Iceland (Cromwell *et al.* 2013), even when accounting for a $\sim 1.3^\circ$ (Model G; Doubrovine *et al.* 2019) or $\sim 0.4^\circ$ (Model TK03; Tauxe & Kent 2004) increase in dispersion from 64°N (northern Iceland) to 71°N (Jan Mayen). Our results may therefore contradict former conclusions (Johnson *et al.* 2008) of a more stable field during the Brunhes.

For the Gilbert reverse chron data, we find that the VGP dispersion, regardless of VGP cut-off, is generally higher than for the Gauss and Matuyama chrons. This may reflect real geomagnetic field behaviour with a more variable field during the Gilbert period. A similar tendency was observed by Kristjansson (2013).

Numerical models (Hatakeyama & Kono 2002) and field observations (Johnson *et al.* 2008) show that the dispersion in VGP positions is expected to vary according to latitude, with a general tendency to increase towards the poles. Several PSV models have been parametrized in terms of expected latitudinal variations in such observations (Johnson *et al.* 2008; Johnson & McFadden 2015). In Fig. 10, we compare our VGP dispersion results (Table 4) against the two most commonly discussed types of models, Model G of McFadden *et al.* (1988) and the TK03 model of Tauxe & Kent (2004), which are based on PSV from recent lavas. We use the revised versions of Model G (Doubrovine *et al.* 2019) for the 0–5 Ma and 0–10

Table 4. VGP dispersion for Fljótisdalur & Jökuldalur data sets combined.

Name	Age interval	Polarity	VGP latitude cut-off	n	S_{mean}	$\frac{S_{\text{upper}}}{S_{\text{lower}}}$	Site mean latitude
$S_{B(1-7)}$	1–7 Ma	N/R	–	154	27.2	$\frac{30.3}{23.9}$	64.6°N
$S_{B(1-7)}$	1–7 Ma	N/R	Van ($>\pm 47^\circ$)	142	21.1	$\frac{22.6}{19.5}$	64.6°N
$S_{B(1-7)}$	1–7 Ma	N/R	$>\pm 45^\circ$	142	21.1	$\frac{22.5}{19.5}$	64.4°N
$S_{B(1-7(N))}$	1–7 Ma	N	–	61	30.3	$\frac{37.1}{23.4}$	64.4°N
$S_{B(1-7(N))}$	1–7 Ma	N	Van ($>\pm 49.6^\circ$)	54	19.7	$\frac{21.9}{17.3}$	64.4°N
$S_{B(1-7(N))}$	1–7 Ma	N	$>\pm 45^\circ$	54	19.7	$\frac{22.1}{17.4}$	64.5°N
$S_{B(1-7(R))}$	1–7 Ma	R	–	93	25.0	$\frac{28.1}{22.0}$	64.7°N
$S_{B(1-7(R))}$	1–7 Ma	R	Van ($>\pm 45.4^\circ$)	88	22.0	$\frac{24.0}{20.0}$	64.7°N
$S_{B(1-7(R))}$	1–7 Ma	R	$>\pm 45^\circ$	88	22.0	$\frac{23.8}{19.9}$	64.7°N
$S_{B(\text{Mat})}$	Matuyama	N/R	–	55	19.8	$\frac{21.7}{17.8}$	65.0°N
$S_{B(\text{Mat})}$	Matuyama	N/R	Van ($>\pm 51.1^\circ$)	54	19.1	$\frac{20.6}{17.3}$	65.0°N
$S_{B(\text{Mat})}$	Matuyama	N/R	$>\pm 45^\circ$	55	19.8	$\frac{21.8}{17.7}$	65.0°N
$S_{B(\text{Gauss})}$	Gauss	N/R	–	23	25.9	$\frac{32.3}{19.5}$	64.4°N
$S_{B(\text{Gauss})}$	Gauss	N/R	Van ($>\pm 42.0^\circ$)	22	23.9	$\frac{29.1}{18.2}$	64.4°N
$S_{B(\text{Gauss})}$	Gauss	N/R	$>\pm 45^\circ$	20	20.4	$\frac{24.2}{15.7}$	64.4°N
$S_{B(\text{Gil})}$	Gilbert	N/R	–	65	31.5	$\frac{37.3}{26.1}$	64.3°N
$S_{B(\text{Gil})}$	Gilbert	N/R	Van ($>\pm 43.2^\circ$)	58	23.2	$\frac{25.5}{20.4}$	64.3°N
$S_{B(\text{Gil})}$	Gilbert	N/R	$>\pm 45^\circ$	58	23.2	$\frac{25.6}{20.8}$	64.3°N

Notes: n, number of independent sites. S_{mean} , $\frac{S_{\text{upper}}}{S_{\text{lower}}}$, mean dispersion with 95% bootstrapped upper and lower confidence levels (Tauxe 2010). Van, latitude cut-off using the Vandamme (1994) criteria.

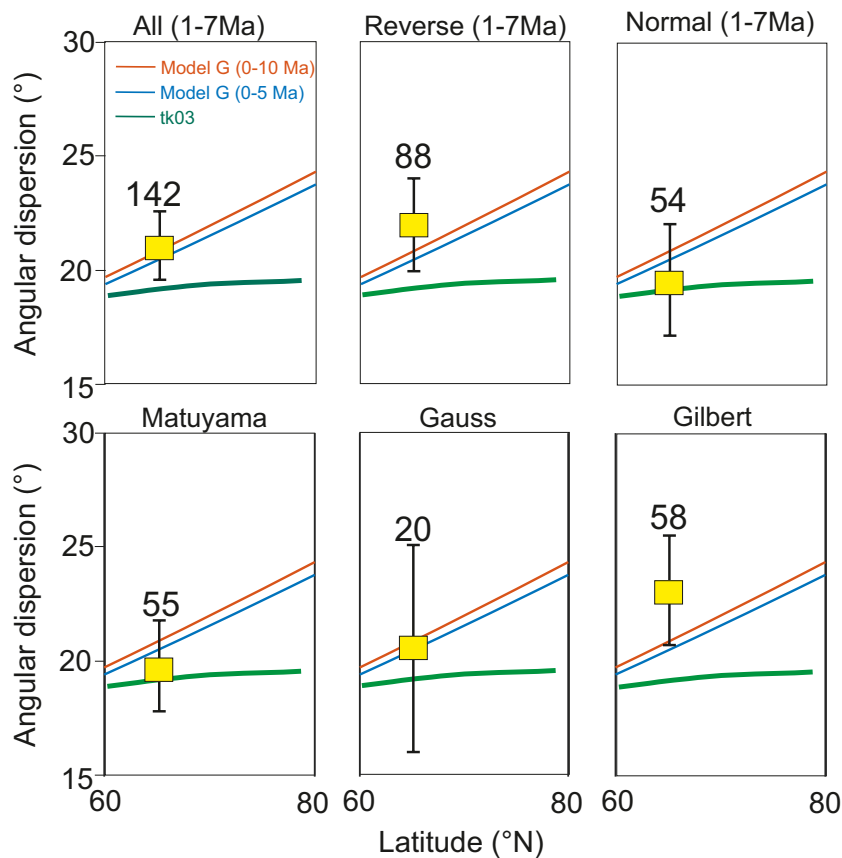


Figure 10. VGP dispersion data for the Fljótisdalur and Jökuldalur palaeodirections using a VGP latitude cut-off of 45° . Upper panels: 1–7 Ma data sets. Lower panels: Matuyama, Gauss and Gilbert polarity chrons. All dispersion values are calculated from data sets with the quality selection criteria $k > 50$ and $N > 5$. Data are shown with bootstrapped 95 per cent levels of confidence error bars. Blue/red lines: revised Model G curves for 0–5 Ma and 0–10 Ma data compilations, respectively (Dobrovine *et al.* 2019). Green line: TK03 (Tauxe & Kent 2004).

Ma intervals based on the recent compilations of (Opdyke *et al.* 2015) and Cromwell *et al.* (2018).

It is clear from Fig. 10 that neither Model G nor TK03 fits the dispersion data consistently. Using all data (top panels in Fig. 10), the data tend to fit the Model G best, except for the 1–7 Ma normal polarity data ($S_{B(1-7(N))}$), which is closer to TK03. Only minor difference is observed for Model G when comparing the 0–5 Ma and 0–10 Ma intervals, which (as also noted by Doubrovine *et al.* (2019)) possibly relates to the fact that most of the data of Cromwell *et al.* (2018) fall within the 0–5 Ma interval. A very nice fit is obtained for all data combined $S_{B(1-7)}$ to Model G, whereas $S_{B(1-7)}$ is not fitted by TK03, even when considering the uncertainty of $S_{B(1-7)}$.

When considering the individual chrons represented by the data set (bottom panels in Fig. 10), we find likewise an inconsistent fit to the two models. While the Matuyama data appear to be closer to TK03, the normal polarity Gauss chron is closer to Model G (although the Gauss angular dispersion is less well constrained due to only 20 data points), and the reverse Gilbert chron data is significantly higher than even the Model G prediction. Considering the uncertainties of the S_b values, all estimates, except that for the Gilbert subset, are consistent with Model G, while the discrepancy with TK03 is generally larger.

6 CONCLUSIONS

In this paper, we present new high-quality palaeodirectional data based on ~1300 cores from 123 accepted sites along five sections in the Fljótsdalur valley in eastern Iceland, at about 65°N. The sites are dominantly found to be of late Miocene to late Pliocene age (2.08–7.00 Ma) based on 17 new high-resolution incremental heating $^{40}\text{Ar}/^{39}\text{Ar}$ age determinations. Only one site is of Brunhes age.

Due to excellent weather conditions during the field campaigns, 96 per cent of all cores were oriented using both a Brunton compass and a sun compass, thus providing us with a unique data set for comparing magnetic and sun azimuths at high northern latitudes. We find that 42, 16 and 3 per cent of the magnetic azimuth estimates (corrected to true north) deviate more than $\pm 5^\circ$, $\pm 10^\circ$ and $\pm 20^\circ$, respectively, from the corresponding sun azimuth estimates. This suggest that core sampling intended for high-quality PSV studies at high northern latitudes should preferably be oriented by sun (or alternatively by a differential GPS).

Based on AF- and thermal-demagnetization experiments, we obtain well-grouped ChRM directions from 113 distinct sites ($N_{\text{mean}} = 9.5$, $k_{\text{mean}} = 353$, $\alpha 95_{\text{mean}} = 3.1^\circ$). It is shown that the mean direction of the data is consistent with the direction of the GAD field and the mean palaeomagnetic pole is not significantly different from the present axis of rotation.

By setting strict quality criteria ($N > 5$, $k > 50$) and combining the Fljótsdalur data with overall 1–2.5 Ma data from Jökuldalur (Døssing *et al.* 2016), we obtain a high-quality data set with 154 independent site mean directions for PSV analysis for the 1–7 Ma time interval. We use these data to calculate dispersion of VGPs and find indications of a higher VGP dispersion during reverse polarity intervals than during normal polarity intervals. We compare individual polarity chrons and find that a higher dispersion tends to be observed for the Gilbert reverse polarity chron as compared to the Gauss and Matuyama chrons. We further compare our results against Brunhes age VGP dispersion data obtained from Jan Mayen (Cromwell *et al.* 2013) and find no evidence of a higher VGP scatter

during the Matuyama as has previously been suggested (Johnson *et al.* 2008).

Importantly, the VGP scatter data S_b presented in this study are generally not well-fitted by the two commonly used models for VGP dispersion: Model G and TK03. However, considering the uncertainties of the S_b values, all estimates, except that for the Gilbert subset, are consistent with Model G, while the Discrepancy with TK03 is generally larger. A very nice fit is obtained for all data combined to Model G, whereas $S_{B(1-7)}$ is not fitted by TK03, even when considering the uncertainty of $S_{B(1-7)}$.

All data presented in this study and by Døssing *et al.* (2016) will be made available to the MagIC database.

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SUPPORTING INFORMATION

Supplementary data are available at [GJI](https://doi.org/10.1017/gji.2024.11) online.

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