

A multiband map of the natural night sky brightness including *Gaia* and *Hipparcos* integrated starlight

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ABSTRACT

The natural night sky brightness is a relevant input for monitoring the light pollution evolution at observatory sites, by subtracting it from the overall sky brightness determined by direct measurements. It is also instrumental for assessing the expected darkness of the pristine night skies. The natural brightness of the night sky is determined by the sum of the spectral radiances coming from astrophysical sources, including zodiacal light, and the atmospheric airglow. The resulting radiance is modified by absorption and scattering before it reaches the observer. Therefore, the natural night sky brightness is a function of the location, time, and atmospheric conditions. We present in this work the GAia Map of the Brightness Of the Natural Sky (GAMBONS), a model to map the natural night brightness of the sky in cloudless and moonless nights. Unlike previous maps, GAMBONS is based on the extra-atmospheric star radiance obtained from the *Gaia* catalogue. The *Gaia*-Data Release 2 (DR2) archive compiles astrometric and photometric information for more than 1.6 billion stars up to G = 21 mag. For the brightest stars, not included in *Gaia*-DR2, we have used the *Hipparcos* catalogue instead. After adding up to the star radiance the contributions of the diffuse galactic and extragalactic light, zodiacal light and airglow, and taking into account the effects of atmospheric attenuation and scattering, the radiance detected by ground-based observers can be estimated. This methodology can be applied to any photometric band, if appropriate transformations from the *Gaia* bands are available. In particular, we present the expected sky brightness for V (Johnson), and visual photopic and scotopic passbands.

Key words: radiative transfer – scattering – atmospheric effects – instrumentation: photometers – light pollution – site testing.

1 INTRODUCTION

The natural brightness of the night sky in different photometric passbands is a relevant parameter for site characterization and light pollution research, since it allows establishing a baseline against which to evaluate the light pollution levels experienced in urban, peri-urban, rural, and pristine dark sites.

Brightness, in this context, is a shortcut name for the spectral radiance of the sky, weighted by the sensitivity function of the photometric band, spectrally integrated over all wavelengths, and angularly integrated within the field of view of the detector.

The measurement of the night sky brightness and the atmospheric conditions, in particular at sites of astronomical interest, has a long history that continues to the present day (e.g. Turnrose 1974 for Palomar and Mount Wilson Observatories; Benn & Ellison 1998 for La Palma; or Patat 2008 for Cerro Paranal). The natural sources of the night sky brightness were already outlined in Roach & Gordon (1973). The reference paper by Leinert et al. (1998) studies and analyses the different sources contributing to the sky brightness covering wavelengths from $\lambda \approx 100$ nm (ultraviolet domain) to $\lambda \approx 200$ µm (far-infrared). It includes airglow, zodiacal light, integrated starlight (ISL), diffuse galactic light (DGL), and extragalactic

background light (EBL), and summarizes the best estimations at the time for all of them. A similar work can be found in Noll et al. (2012) for the wavelength range from $\lambda \approx 300$ to 920 nm, with an extended discussion on the airglow and atmospheric extinction at Cerro Paranal. Duriscoe (2013) evaluates the artificial contribution to the sky brightness by subtracting the natural contributions from the overall sky radiance measured with all-sky imaging systems. Duriscoe (2013) used a model of ISL (including DGL) constructed from images made with the same equipment used for sky brightness observations, and modelling zodiacal light and airglow from previously available data.

In parallel, several models to compute the artificial light contribution to the overall night sky brightness have been developed, from the early works by Treanor (1973), Berry (1976), and Garstang (1986, 1989, 1991) to the more recent by Aubé et al. (2005, 2018), Kocifaj (2007, 2018), and Cinzano & Falchi (2012). The outcome of these models are maps of the artificial brightness of the night sky (for instance Cinzano, Falchi & Elvidge 2001; Falchi et al. 2016) to which the natural brightness shall be added in order to obtain the total one, which is the physical observable available from direct measurements.

The publication of the *Gaia*-Data Release 2 (DR2) archive (Gaia Collaboration et al. 2018), with high-quality space-based global astrometric and photometric information for more than 1.6 billion

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stars in all the sky up to G = 21 mag (G being the white light magnitude observed with Gaia in the astrometric instrument), offers a unique opportunity to improve the computation of the ISL, by directly integrating the contributions of all the stars in the Gaia catalogue. As a first result, we present in this work the first maps of the extraatmospheric brightness of the night sky in the V (Johnson) and the human-vision photopic and scotopic bands, directly obtained from the *Gaia* and *Hipparcos* (ESA 1997; van Leeuwen 2007) catalogues. Hipparcos is used here to extend Gaia catalogue to the bright end, not included in Gaia-DR2. The inclusion of the remaining relevant sources, including zodiacal light and airglow, and the corrections due to attenuation and scattering into the field of view allow to compute the radiance at ground level at any direction of the sky and any location and night time for the observer. This methodology can be applied to any photometric band, if appropriate transformations from the native Gaia bands are computed. In particular, we have determined the sky brightness for the Johnson V and human-vision photopic and scotopic passbands. All together constitutes the first version of the GAia Map of the Brightness Of the Natural Sky (GAMBONS) model of the natural night sky brightness.

The novelty of this work is the determination of the radiance outside the Earth's atmosphere using *Gaia* and *Hipparcos* catalogues. However, a complete sky brightness model requires the modelling of the other components, as zodiacal light, airglow, or atmospheric attenuation and scattering. We have implemented in GAMBONS the more updated and reliable models available in the literature for those components, as described in the next sections, but it is not the scope of this work to discuss in detail each one of them. On the other hand, the final map of the night sky brightness is highly dependent on the atmospheric conditions, including airglow. Thus, the accuracy of the results depends on the reliability of the atmospheric parameters used in the computation.

The *Gaia* and *Hipparcos* catalogues are overviewed in Section 2. Their photometric systems and the other photometric bands used in this work are described in Section 3. Section 3 also deals with the problem of deriving the needed transformations between different photometric bands. Section 4 describes the general characteristics of our model. Sections 5–7 explain in detail the computation of the different sources of the natural night sky brightness and the models for atmospheric attenuation and scattering. Finally, we expose some applications of the model in Section 8.

2 THE GAIA AND HIPPARCOS CATALOGUES

Gaia (Gaia Collaboration et al. 2016) is an extremely ambitious astrometric space mission of the scientific programme of the European Space Agency (ESA). Gaia measures, with very high accuracy, the positions, motions, and parallaxes of a large number of stars and many other kinds of objects as galaxies, asteroids, and quasars.

Consequently, a detailed three-dimensional map of more than 1.6 billion stars of our Galaxy (approximately 1 per cent of the stars populating the Milky Way) is obtained. This map also includes, for almost all the objects, information about their brightness and colour, as well as radial velocity, and several astrophysical parameters for a large fraction of them.

Gaia-DR2 photometry in G band is almost complete up to $G \approx 20$ mag, with a limiting magnitude slightly fainter than G = 21 mag in some areas of the sky. Gaia extends to the faint the catalogue obtained by the Hipparcos ESA mission, launched in 1989. Hipparcos catalogue includes information about the position, motion, brightness, and colour of the sources. Although the Gaia catalogue is several orders of magnitude more complete and precise

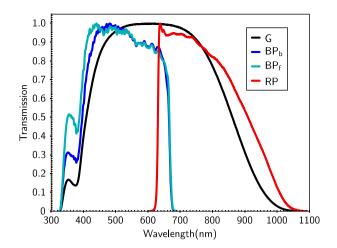


Figure 1. *Gaia* passbands transmission used in this work (Maiz Apellaniz & Weiler 2018).

than Hipparcos catalogue, Gaia cannot observe very bright objects (with G < 5 mag). For those stars, the Hipparcos catalogue addresses the lack of Gaia data. In particular, of the 118 218 stars in the Hipparcos catalogue (ESA 1997), 83 034 are present in the Gaia-DR2-Hipparcos best neighbour catalogue available on the Gaia archive, and has therefore, Gaia data available. The details of the Gaia-DR2-Hipparcos cross-match are given in Marrese et al. (2019). For the 35 000 remaining stars not having Gaia information, we have used photometry from the Hipparcos catalogue. In terms of flux, the Hipparcos stars account for around 20 per cent of the total ISL. After these considerations, our final catalogue contains 1692 953 899 stars.

3 THE PHOTOMETRIC DATA

3.1 The Gaia photometric system

The Gaia broad-band photometric system is fully described in Jordi et al. (2010). It is composed of three different passbands (G, $G_{\rm BP}$, and G_{RP}). The transmission curves represented in Fig. 1 were derived by Maiz Apellaniz & Weiler (2018), as the most representative of the Gaia-DR2 data according to the Gaia team. G is a panchromatic band covering the wavelength range from about 350 to 1000 nm. It is the result of the transmission of the mirrors and the response of the instruments in Gaia. $G_{\rm BP}$ (blue photometer) and $G_{\rm RP}$ (red photometer) are two broad passbands in the wavelength ranges 330-680 and 640–1000 nm, respectively. The corresponding magnitudes are obtained from the integrated flux of the low-resolution spectra of the Gaia spectrophotometer. These photometric data allow the classification of the sources by deriving the astrophysical parameters, such as effective temperature, gravity, chemical composition, and interstellar absorption. An additional description of the photometric contents of Gaia-DR2, including uncertainty distribution, external comparisons, and colour transformations, can be found in Evans et al. (2018).

A small systematic inconsistency in the BP photometric system was spotted by Maiz Apellaniz & Weiler (2018) and Weiler (2018), likely caused by insufficient convergence of the $G_{\rm BP}$ calibration in Gaia-DR2 for bright sources. Maiz Apellaniz & Weiler (2018)

¹https://www.cosmos.esa.int/web/gaia/dr2-known-issues

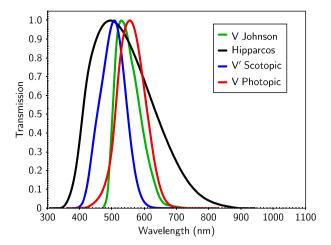


Figure 2. V Johnson (Bessell & Murphy 2012), Hipparcos (H_p) (Weiler et al. 2018), V scotopic (CIE 1951), and V photopic (CIE 1990) passbands transmission used in this work.

propose to mitigate this effect by using two different $G_{\rm BP}$ response curves ($G_{\rm BP}^b$ and $G_{\rm BP}^f$) for the magnitude ranges brighter and fainter than G=10.87 mag, respectively.

3.2 Photometry from Hipparcos catalogue

Hipparcos catalogue provides photometric magnitudes in several passbands. Three of them, H_p , B_T , and V_T , are the three native Hipparcos-Tycho bands. Furthermore, the catalogue compiles external values of the V Johnson magnitude and (B-V) and (V-I) colours in the Johnson-Cousins system (Johnson 1963).

In this work, we use V and H_p magnitudes and (B-V) and (V-I) colours to transform to other passbands, as described in Section 3.4. The passband transmissions used here for V (Bessell & Murphy 2012) and H_p (Weiler et al. 2018) are shown in Fig. 2. In the case of the V band, the *photonic response function* defined in Bessell & Murphy (2012) was transformed back to its original energy response form, since the classical V magnitude system is defined in terms of in-band irradiances, not photon numbers.

3.3 Scotopic and photopic visual passbands

The scotopic $V(\lambda)$ and photopic $V(\lambda)$ spectral responses (Fig. 2) describe the response of the human eye fully adapted to low (<0.005 cd m⁻²) and high (>5 cd m⁻²) environmental luminances, respectively. These photometric bands have been standardized by CIE (1951, 1990). The luminance of a radiance field seen with photopic adaptation, in cd m⁻², is obtained by multiplying the integrated radiance within the $V(\lambda)$ band (in W m⁻² sr⁻¹) by the photopic luminous efficacy factor $K_{\rm m}=683~{\rm lm}~{\rm W}^{-1}$. Similarly, the luminance of a radiance field seen under scotopic adaptation is given by the radiance in the $V(\lambda)$ band multiplied by the scotopic luminous efficacy factor $K'_{\rm m} = 1700~{\rm lm}~{\rm W}^{-1}$. Mesopic sensitivity bands can also be defined for intermediate adaptation states (0.005-5 cd m⁻²; CIE 2010). The luminance is the basic physical quantity at the root of the human perception of brightness. Note that the 'visible' V Johnson band does not provide an accurate estimation of the visual brightness of the sky: skies with the same brightness in V Johnson may correspond to very different luminances, depending on the spectral composition of the sky radiance (Bará et al. 2020), and, of course, on the adaptation state of the observer.

3.4 Transformations between passbands

The purpose of this work is to obtain the maps of the natural night sky radiance in any photometric band, assuming that its spectral response is known. The input data available from *Gaia* and *Hipparcos* catalogues are not sky radiances (W m⁻² sr⁻¹), but the G, G_{BP} , G_{RP} , V magnitudes, and the (B - V) and (V - I) colours of the individual stars. Therefore, we need to transform from these star magnitudes and colours in their native bands to the corresponding sky radiance in the desired target band. Let us recall that astronomical magnitudes are a logarithmic measure of the in-band irradiances (W m⁻²) at the entrance pupil of the observing instrument. The canonical transformation procedure then consists of three main steps: (i) in the first one, the catalogue magnitudes are used to compute the irradiances produced by each individual star in the original bands; (ii) the star irradiances in the original bands are used to estimate the star irradiances within the target observation band; and finally (iii) the radiance L of any patch of the sky in the target band is given by the sum of the irradiances contributed by all stars contained within that patch, divided by the solid angle (in steradian) it subtends (Bará et al. 2020). This radiance is usually called surface brightness in the astrophysical literature.

In practice these steps can be performed in different ways, and in some cases they can be grouped into a single transformation. The details depend on the available data. Note, for instance, that V Johnson is included in the Hipparcos catalogue but not in Gaia-DR2. So, we need to transform the Gaia data from G to V. Furthermore, the input provided by the Gaia-DR2 catalogue is given also in photons per second within the instrument entrance pupil, which is a scaled version of the actual in-band photon irradiance (photons s^{-1} m⁻²) and hence of the irradiance (W m⁻²), being the details of this last transformation contingent on the stars' spectra.

In general, a transformation between two passbands (A and B) is a function of, at least, one colour. For many cases, a polynomial expression with only one colour C, derived by subtracting the magnitude in two different passbands, is enough to obtain a good fit between A and B, in the form

$$A = B + \sum_{i} a_i C^i. \tag{1}$$

Taking into account the relation $A = A_0 - 2.5 \log (E_A/E_0)$, between the magnitude A and the corresponding irradiance E_A , with A_0 the zero-point magnitude for the E_0 irradiance, we can get a relation for the ratio E_A/E_B from equation (1):

$$\frac{E_A}{E_B} = e^{\sum_i a_i C^i}.$$
 (2)

The A_0 and B_0 constants and the conversion from common to natural logarithm are included in the e^{a_0} term.

For instance, the transformation between G and V can be expressed as function of the colour $(G_{\rm BP}-G_{\rm RP})$:

$$E_V = E_G e^{\sum_i a_i (G_{\rm BP} - G_{\rm RP})^i}.$$
 (3)

The methodology to get the coefficients a_i of the transformations is the same as in Jordi et al. (2010). First, the set of stellar spectra in BaSeL-3.1 models (Westera et al. 2002) is used to compute the synthetic photometry in all the bands. BaSeL-3.1 covers a wide range of effective temperature, surface gravity, and stellar metallicity including all possible evolutionary stages in the stellar evolution. We

also incorporate the effect of interstellar reddening by using a wide range of possible absorption rates (from 0 to 11 mag at 550 nm) of the *Gaia* and *Hipparcos* stars. Following Jordi et al. (2010), a value equal to 0.03 mag has been assumed for each synthetic magnitude of a Vega-like star. The synthetic photometry was simulated in both photons per second within the *Gaia* input pupil and in W m⁻², allowing the transformation of units to be included in the fit (for instance from the radiant flux F_G in photons per second within the *Gaia* input pupil to the corresponding irradiance E_V in W m⁻²; or from the irradiance E_V to the scotopic irradiance E_{Sco} , both in W m⁻²). Finally, a minimum least-squares fit is applied to get the a_i coefficients in equation (2).

The conversion from the catalogue astronomical magnitude m of each star to the irradiance it produces, E (W m⁻²), is made using the standard definition of magnitudes.

$$E = E_{\rm r} \, 10^{-0.4m},\tag{4}$$

where $E_{\rm r}$ is the reference irradiance associated with the 'zero-point' of the band (V or $H_{\rm p}$ for *Hipparcos* stars), i.e. the irradiance corresponding to m=0.0 mag, given by the integral

$$E_{\rm r} = \int_0^\infty S(\lambda) E_{\rm Vega}(\lambda) \, \mathrm{d}\lambda,\tag{5}$$

where $E_{\text{Vega}}(\lambda)$ is the STIS003 spectral irradiance of Vega from Bohlin & Gilliland (2004), normalized in such a way that $m_{\text{Vega}} = 0.03$ mag for all the bands, and $S(\lambda)$ is the photometric passband.

As pointed out above, the radiance L of any region of the sky is computed as the sum of the irradiances produced by all sources contained in that region, divided by the region's angular extent in steradians. This radiance can be expressed in the logarithmic scale of magnitudes per square arcsecond within the $S(\lambda)$ band, m_S , according to the definition

$$L = L_{\rm r} \, 10^{-0.4 m_{\rm S}},\tag{6}$$

where $L_{\rm r}$ is the reference radiance given by

$$L_{\rm r} = \frac{1}{K} \int_0^\infty S(\lambda) E_{\rm Vega}(\lambda) \, \mathrm{d}\lambda,\tag{7}$$

being $K = 2.3504 \times 10^{-11}$, the steradian equivalent of 1 arcsec².

In principle, this transformation methodology can be applied to any band, taking care to choose the colour that minimizes the residuals of the fit. For *Gaia* passbands this colour is $(G_{\rm BP}-G_{\rm RP})$, while for V, $H_{\rm P}$, and human-vision bands we choose (V-I) or (B-V), depending on the transformation. The relationships between the different bands used in this work are shown in Fig. 3. The a_i coefficients, the colours C, and the σ of the fits for all of them are summarized in Table 1. The values of $L_{\rm T}$ for several bands used in this work are given in Table 2.

4 THE MODEL

We describe in this section the general expression for the inband integrated radiance, $L_{\rm obs}(\alpha,h)$, detected by an observer in the direction $\alpha=(a,z)$ (a the azimuth and z the zenith angle) relative to its reference frame, at height h above sea level. $L_{\rm obs}(\alpha,h)$ must take into account all the sources contributing to the natural night sky brightness: the ISL; the DGL; the EBL and the zodiacal light, which conform the radiance outside the Earth's atmosphere; and the atmospheric airglow. This extra-atmospheric radiance $L_0(\lambda,\alpha)$ (formally including airglow) contributes to the overall map of the sky as seen by the observer as follows:

(i) it produces a direct (attenuated) radiance $L_d(\lambda, \alpha, h) = L_0(\lambda, \alpha) T(\lambda, z, h)$ in the direction of observation α ;

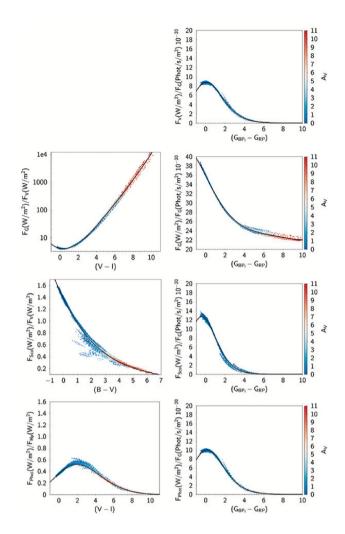


Figure 3. Relationship between the irradiances at different bands used in this work obtained using BaSeL-3.1 spectral library (Westera et al. 2002). The colour code represents the interstellar absorption (vertical scale). The black line is the fit of equation (2) with the coefficients in Table 1.

(ii) it introduces a scattered radiance $L_s(\lambda, \alpha_s, \alpha, h)$ in the remaining pixels of the sky, including the pixel $\alpha = \alpha_s$ itself.

 $T(\lambda, z, h)$ is the effective atmospheric transmittance at wavelength λ , accounting for the radiance reduction due to the attenuation along the beam path from the limits of the atmosphere to the location of the observer.

If we denote $L_s(\lambda, \alpha, h)$ as the total scattered radiance reaching the observer when looking at the direction α ,

$$L_{\text{obs}}(\lambda, \alpha, h) = L_{\text{d}}(\lambda, \alpha, h) + L_{\text{s}}(\lambda, \alpha, h), \tag{8}$$

$$L_{\text{obs}}(\boldsymbol{\alpha}, h) = \int_{0}^{\infty} S(\lambda) L_{\text{obs}}(\lambda, \boldsymbol{\alpha}, h) \, d\lambda. \tag{9}$$

The effects of the atmospheric transmittance and scattering, i.e. the computation of $T(\lambda, z, h)$ and $L_s(\lambda, \alpha, h)$, are described in detail in Section 7.

 $L_{\rm obs}(\lambda, \alpha, h)$ can be expressed as the sum of the radiance coming from different contributors: $L_{\rm obs,ISL}(\lambda, \alpha, h)$ (ISL); $L_{\rm obs,DGL}(\lambda, \alpha, h)$ (DGL); $L_{\rm obs,EBL}(\lambda, \alpha, h)$ (EBL); $L_{\rm obs,zl}(\lambda, \alpha, h)$ (zodiacal light); and

Table 1. Coefficients and mean standard deviations of the transformations between irradiances in different bands, according to equation (2). $E^{'}$ denotes irradiances in photon s⁻¹ m⁻², whereas E denotes irradiance in W m⁻². For the transformations from Gaia photometry, the last column shows the standard deviation for the range $(G_{BP} - G_{RP}) < 3$ mag, where the most of the Gaia-DR2 stars are.

	С	a_0	a_1	a_2	a_3	a_4	a_5	$\sigma_{\rm all} (\%)$	σ _{Gaia} (%)
E_G/E_G'	$G_{ m BP}^{ m f}-G_{ m RP}$	- 42.474	- 0.1168	- 0.006034	+0.005261	-6.496×10^{-4}	$+2.521 \times 10^{-5}$	0.5	0.4
E_V/E_G'	$G_{ m BP}^{ m f'}-G_{ m RP}$	-43.878	+0.028640	-0.20573	+0.017942	-5.444×10^{-4}		5.1	2.4
$E_{\rm Sco}/E_G'$	$G_{ m BP}^{ m f'}-G_{ m RP}$	-43.551	-0.201	-0.2284	+0.02308	-8.216×10^{-4}		12.5	6.0
E_{Phot}/E_G'	$G_{\mathrm{RP}}^{\mathrm{f}}-G_{\mathrm{RP}}$	-43.749	+0.04745	-0.1972	+0.01678	-5.03×10^{-4}		3.8	2.2
E_G/E_V	(V-I)	1.4221	-0.14355	+0.22442	-0.020977	$+7.827 \times 10^{-4}$		4.3	
$E_{\rm Sco}/E_V$	(B-V)	0.31760	-0.34836	-0.017739	+0.0046171	-4.0399×10^{-4}		5.1	
$E_{\text{Phot}}/E_{H_{\text{p}}}$	(V-I)	-1.0155	+0.39292	-0.12037	+0.0077708	-3.2423×10^{-4}		2.7	

Table 2. Reference radiances $L_{\rm T}$ based on STIS003 spectrum and $G=V=H_{\rm p}=0.03$ mag for Vega.

Passband	$({\rm Wm^{-2}sr^{-1}})$
\overline{G}	562.5373
V	143.1685
$H_{ m p}$	426.8769

 $L_{\text{obs,ag}}(\lambda, \boldsymbol{\alpha}, h)$ (airglow):

$$\begin{split} L_{\text{obs}}(\boldsymbol{\alpha},h) &= \sum_{\forall \mathcal{P}} \left[\int_{0}^{\infty} S(\lambda) L_{\text{obs},\mathcal{P}}(\lambda,\boldsymbol{\alpha},h) \, \mathrm{d}\lambda \right] \\ &= L_{\text{obs,ISL}}(\boldsymbol{\alpha},h) + L_{\text{obs,DGL}}(\boldsymbol{\alpha},h) \\ &+ L_{\text{obs,EBL}}(\boldsymbol{\alpha},h) + L_{\text{obs,zl}}(\boldsymbol{\alpha},h) + L_{\text{obs,ag}}(\boldsymbol{\alpha},h). \end{split}$$

$$(10)$$

In the next sections we will first describe in detail the computation of the extra-atmospheric radiance for each component, and thereafter the effect of the atmospheric attenuation and scattering.

Note that the different components of the radiance are usually given in different coordinate systems: equatorial (α, δ) or galactic (l, b) coordinates for the astrophysical component; ecliptic (Λ, β) for the zodiacal light; and horizontal (a, \hat{h}) coordinates for the airglow and atmospheric attenuation, \hat{h} being the angular height above the horizon. Horizontal coordinates are linked to the local reference frame of the observer and therefore the horizontal coordinates of any extraterrestrial source depend on the observer position and time. Moreover, both the airglow $L_{\rm ag}$ and the atmospheric transmittance T depend only on the zenith angle $z=90-\hat{h}$. For the details about coordinate transformation, see for instance Green (1985).

5 RADIANCE OUTSIDE THE EARTH ATMOSPHERE

5.1 Integrated starlight

In the visual wavelength range, the ISL is one of the most important contributors to the natural sky brightness. In some circumstances, as for lines of sight towards the Galactic Centre far from the Sun and for low solar activity, ISL can be the brightest component. Early works on the night sky brightness used simple models of the Galaxy to determine the ISL. Bahcall & Soneira (1980) constructed a model with the Galaxy consisting of an exponential disc and a power-law spheroidal bulge that was widely used in the last years of the last century.

Other early approach to the problem of the ISL was to use data from Imaging Photopolarimeters (IPPs) on the *Pioneer 10* and *Pioneer 11* deep space probes (Weinberg et al. 1974). The full set of data contains stars and background integrated light for almost all the sky with a spatial resolution around 2° in two bands, blue and red.

From the early 2000s, the determination of the ISL takes advantage of the emergence of large surveys, as Tycho-2, United States Naval Observatory A2 (USNO-A2), or Two Micron All-Sky Survey (2MASS; see e.g. Melchior, Combes & Gould 2007). Following this line, the publication of *Gaia*-DR2 offers a new opportunity to determine the ISL from the most complete photometric survey ever published.

Before starting with the description of the use of the *Gaia* data to establish the ISL, it is also worth mentioning the work of Duriscoe (2013). He used more than 700 sky images taken from pristine mountaintop locations to get the contribution of the stellar contribution plus diffuse galactic and extragalactic light, after removing the contribution from airglow and zodiacal light.

5.1.1 ISL from Gaia-DR2 and Hipparcos

With the transformations described in Section 3.4, it is possible to calculate the irradiance produced at the top of the atmosphere in a given passband by each and every star in our catalogues, and then to compute the sky map of the ISL by first (a) adding these irradiances within small elementary patches of the sky, and (b) converting them into radiances by dividing the irradiances by the solid angle (in steradian) subtended by each patch. For that, the sky is first tessellated into N pixels by using the Hierarchical Equal Area isoLatitude Pixelisation (HEALPix) scheme (Górski et al. 2005). We used a resolution equal to 8 and therefore $N=786\,432$ pixels of mean area $0.05245\,\deg^2$ (equivalent to $1.5979\times10^{-5}\,\mathrm{sr}$), but the methodology could be applied to any desired HEALPix resolution.

For each pixel we have collected all the stars in Gaia-DR2 and Hipparcos catalogues and added its irradiances for all the considered passbands. However, there are more than 300 million of faint sources without colour information in Gaia-DR2. In order to not lose these stars, we have assigned to them the mean $G_{\rm BP}-G_{\rm RP}$ colour of the stars in its surrounded area, allowing in this way the application of the transformation described in Section 3.4. The in-band irradiances E^* for each star are computed using equation (2) with the appropriate coefficients and colour given in Section 3.4:

$$E_p = \sum_{k=1}^{N_p} E_k^*, \tag{11}$$

where E_p stands for irradiance in a given band in the pixel p, and N_p stands for the number of stars inside the pixel. This total irradiance

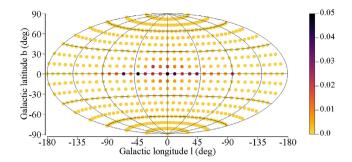


Figure 4. Ratio between flux in photons s^{-1} in the G band of faint stars (20 < G < 27.5 mag) and bright stars (10.5 < G < 21 mag) obtained using the Besançon Galaxy Model.

is divided by the solid angle $\Delta\omega$ subtended by the pixel to obtain the radiance outside the atmosphere from the ISL, $L_{\text{ISL},p}(l,b) = E_p/\Delta\omega$, with (l,b) the galactic coordinates of the centre of the pixel. Following the notation introduced in Section 4, with $L_k^*(\lambda)$ the contribution to the radiance at wavelength λ of the kth star in the pixel, $L_{\text{ISL},p}(l,b)$ can be expressed as

$$L_{\text{ISL},p}(l,b) = \sum_{k=1}^{N_p} \int_0^\infty L_k^*(\lambda) S(\lambda) \, d\lambda. \tag{12}$$

The integral in equation (12) cannot be computed explicitly, as we have not $L_k^*(\lambda)$ for each star. But it is equal to the flux in G from Gaia-DR2 transformed to in-band irradiances expressed in W m⁻², following equation (2), and divided by the pixel solid angle, $\Delta \omega$.

Stars fainter than G = 20.0 mag, the approximated limit of the Gaia-DR2 completeness, could contribute with some amount of radiance to the ISL. We have used the Besançon Galaxy Model (Robin et al. 2003, 2012) in order to estimate this lost flux. The model provides a realistic description of the stellar content of the Milky Way, including its kinematic and dynamics, the mass distribution, the star formation rate, and evolution of different stellar populations. It has been used to simulate the stars in the Milky Way up to G = 27.5 mag in a grid of patches between 0.25 and 1 deg² distributed in a $10^{\circ} \times 10^{\circ}$ grid in (l, b) over the sky. These stars have been classified in two groups: bright stars (10.5 < G < 20.0 mag) and faint stars (21.0 <G < 27.5 mag). The stars with G between 20.0 and 21.0 mag have been assigned to one group and another following a probability linear distribution that approximately reproduce the selection function of the Gaia-DR2 catalogue. Then the ratio between the total flux of both groups was computed (Fig. 4). Only a few points on the Galactic Plane and around the Galactic Centre have a contribution from faint stars >3 per cent. As we show in Section 8, the contribution of the Milky Way to the total sky brightness could reach, under some conditions and in some lines of sight, the most important contributor to the total night sky brightness. So, we decided to correct the total radiance $L_{ISL, p}(l, b)$ by adding the contribution of the very faint stars not present in Gaia-DR2, computed from the ratios obtained from the Besançon Galaxy Model in order to improve the accuracy of our model.

5.2 Background galactic light

The DGL is the diffuse background radiation produced by the scattering of the starlight in the dust grains present in the interstellar space. It contributes typically between 20 and 30 per cent of the total integrated light from the Milky Way (Leinert et al. 1998).

DGL is very difficult to map because it is masked by the light coming from the unresolved stars and, for ground-based observations, airglow and zodiacal light. Despite all this, several works to characterize and even model the DGL in the optical range have been done. A simple estimation of the DGL can be done using the relation between the ISL and DGL intensities in a given line of sight. It is supported by the fact that DGL is mainly originated from the forward scattering on interstellar grains, tracking in this way the starlight in a given sky direction. Leinert et al. (1998) give the ratios DGL to ISL for different galactic latitudes based on Toller (1981). A second and more accurate approach is to use the data from *Pioneer* probes (Arai et al. 2015). The IPP instruments on Pioneer 10 and Pioneer 11 collected data in blue (395–495 nm) and red (590–690 nm) bands. The measured radiances at heliocentric distances greater than 3 au are assumed as not affected by the zodiacal light. After removing the contribution of the ISL using stellar counts or synthetic models, it is possible to get the DGL plus the EBL (the integrated radiation from all light sources outside the Galaxy). A summary of the use and limitations of this methodology can be found in Toller (1990).

In this work, we have adopted a different approach. From the early work of Laureijs, Mattila & Schnur (1987), several authors have pointed out the relation between the diffuse emission of the dust at $100~\mu m$ and its emission at visible wavelength. The relation can be written, using the notation in Matsuoka et al. (2011) and Kawara et al. (2017), as

$$I_{\nu,i}(DGL) = b_i I_{\nu,100} - c_i I_{\nu,100}^2,$$
(13)

$$I_{\nu.100} = I_{\nu.SFD} - 0.8 \,\text{MJy sr}^{-1},$$
 (14)

where $I_{\nu,\,100}$ is the spectral radiance at 100 μm from the interstellar medium (ISM), b_i and c_i are free parameters, and $I_{\nu,\,SFD}$ is the 100 μm spectral radiance from the diffuse emission map of Schlegel, Finkbeiner & Davis (1998, hereafter SFD). Equation (14) accounts for the EBL that must be subtracted from the SFD map. The EBL emission at 100 μm is ≈ 0.8 MJy sr⁻¹ (Matsuoka et al. 2011). Their optical and 100 μm emissions are not correlated.

The negative quadratic term in equation (13) reflects the observed saturation in the DGL radiance when regions with high 100 μ m emission become optically thick (Ienaka et al. 2013). The use of equation (13) is therefore limited to $I_{\nu, SFD} < 50$ MJy sr⁻¹. This restriction applies mainly to low galactic latitudes $|b| \lesssim 30^{\circ}$, where high 100 μ m emission at very optical thick regions is found. In this cases, equation (13) could give preposterous DGL radiances several times greater than ISL ones. To avoid this, we have imposed an upper limit in the DGL to ISL radiance ratio equal to 0.35, according with the highest values reported by Toller (1981).

Our computation of DGL is based on the $100~\mu m$ spectral radiance SFD map. It is a combination of the *Infrared Astronomical Satellite (IRAS)* and the *Cosmic Background Explorer (COBE)*/Diffuse Infrared Background Experiment (DIRBE) data, with a resolution of few arcminute. It is the main source to derive the dust temperature, opacity, and extinction. Values are given in MJy sr⁻¹ (1 MJy = 10^{-20} W m⁻² Hz⁻¹). The zodiacal foreground emission and bright stars have been removed from the map, but not the Large Magellanic Cloud (LMC), Small Magellanic Cloud (SMC), and M31 extragalactic sources.

The coefficients b_i and c_i are taken from Kawara et al. (2017), who give their values for several optical wavelengths from 0.23 to 0.65 μ m. This range of wavelengths almost fully covers the optical bands considered in this work, allowing the interpolation of the b_i and c_i given in Kawara et al. (2017) for the whole range and therefore the computation of the radiance of the DGL, $L_{\rm DGL}(\lambda)$, following the

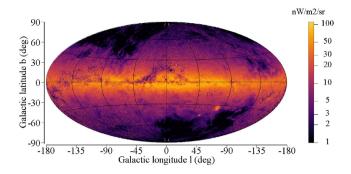


Figure 5. Diffuse galactic light (DGL) radiance in the V Johnson passband.

same steps as for the ISL. The resultant map of the DGL in the V band is shown in Fig. 5.

5.3 Extragalactic background light

The extragalactic background light (EBL) is a minor contributor to the sky brightness. EBL is the sum of all extragalactic sources, mainly resolved and unresolved galaxies and intergalactic matter. It is assumed as isotropic. Because of its low intensity, its observation and measurement are strongly disturbed by the zodiacal light and airglow. In spite of its low contribution, we have included ELB in our model, based on the data from Driver et al. (2016). In that paper, the EBL is derived for a wide range of wavelengths, including the UBVI bands, from a combination of wide and deep galaxy numbercount data from the Galaxy And Mass Assembly (GAMA), Cosmic Evolution Survey (COSMOS)/G10, Hubble Space Telescope (HST) Early Release Science, HST Ultraviolet Hubble Ultra Deep Field (UVUDF), and various near-, mid-, and far-IR data sets from ESO, Spitzer, and Herschel. We used the data of table 2 of that paper to interpolate at any wavelengths in the optical range. With this procedure, we obtain a radiance of the EBL at the V band equal to $1.1 \text{ nW m}^{-2} \text{ sr}^{-1}$.

The final map of the radiance in V Johnson passband outside the atmosphere including the ISL, the DGL, and the EBL is shown in

Fig. 6. These data, together with the radiances for the *G* and scotopic and photopic passbands, are available online. A sample of the data is shown in Table 3.

5.4 Zodiacal light

Zodiacal light is originated by the scattering of the Sun light in the dust particles near the ecliptic plane. It represents a significant term of the natural night sky brightness. The zodiacal light decreases with the angular distance to the Sun, with the exception of the antisolar point, where we find the *gegenschein* contribution caused by the backward scattering of the solar light. It also decreases with the heliocentric distance. The values for the optical emission (at $\lambda = 500$ nm) of the zodiacal light at 1 au were compiled by Levasseur-Regourd & Dumont (1980) and updated by Leinert et al. (1998). Kwon, Hong & Weinberg (2004) present a new set of data, with a smaller sky coverage near the Sun than the previous work. Both sets of data have a high degree of agreement and offer confident values for the zodiacal light brightness. In this work, we have used the data given in Leinert et al. (1998).

The spectrum of the zodiacal light is slightly reddened with respect to the solar spectrum, and therefore a colour correction must be taken in to account when computing its radiance in other bands different to V. Leinert et al. (1998) account for the effect of the colour in the zodiacal light through the factor $f_{\rm co}$, a measure of the quotient of the zodiacal light and the solar radiance, normalized at $\lambda=500$ nm, function of the wavelength and elongation (ϵ):

$$f_{\rm co}(\lambda, \epsilon) = \frac{L_{\rm zl}(\lambda)/L_{\bigodot}(\lambda)}{L_{\rm zl}(500)/L_{\bigodot}(500\,{\rm nm})}.$$
 (15)

From the values of $f_{co}(\lambda, \epsilon)$ in Leinert et al. (1998), the zodiacal light spectrum at 1 au can be obtained from the solar spectrum:

$$L_{\rm zl_0}(\lambda, \Lambda, \beta) = f_{\rm co}(\lambda, \epsilon) L_{\rm zl_0}(500) \frac{{\rm E}_{\odot}(\lambda)}{{\rm E}_{\odot}(500)}, \tag{16}$$

where E_{\odot} is the STIS002 spectral irradiance of the Sun from the CALSPEC library (Bohlin, Gordon & Tremblay 2014), and the elongation can be computed from the ecliptic coordinates (Λ, β) .

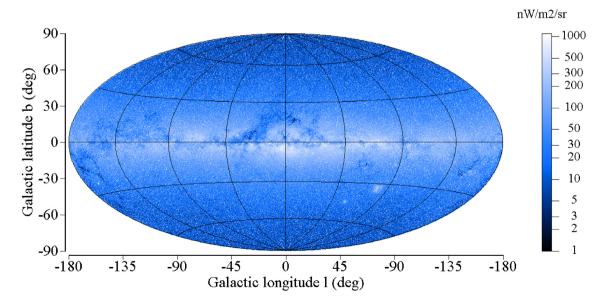


Figure 6. Sky map of the radiance outside the Earth atmosphere (ISL, DGL, and EBL) in the V Johnson passband.

Table 3. Radiances (W m^{-2} sr⁻¹) outside the Earth atmosphere including the integrated starlight (ISL), the diffuse galactic light (DGL), and the extragalactic background light (EBL) for the V Johnson, G, and scotopic and photopic passbands. The full table is available online as supplementary material.

HEALPIX Id	l_{gal} (°)	b_{gal} (°)	L_V	L_G	$L_{ m Sco}$	$L_{ m Phot}$
0	45.000	0.149	6.31994×10^{-8}	5.27586×10^{-7}	5.81792×10^{-8}	1.39854×10^{-7}
1	45.176	0.298	8.43552×10^{-8}	5.78342×10^{-7}	8.25604×10^{-8}	1.57852×10^{-7}
2	44.824	0.298	7.02885×10^{-8}	4.89853×10^{-7}	6.72805×10^{-8}	1.45569×10^{-7}
3	45.000	0.448	9.13240×10^{-8}	7.20925×10^{-7}	8.42789×10^{-8}	1.65076×10^{-7}
4	45.352	0.448	9.30920×10^{-8}	6.45215×10^{-7}	8.92242×10^{-8}	1.65972×10^{-7}
5	45.527	0.597	1.78077×10^{-7}	1.13703×10^{-6}	1.69802×10^{-7}	2.41774×10^{-7}
6	45.176	0.597	7.68996×10^{-8}	5.58962×10^{-7}	7.35909×10^{-8}	1.51565×10^{-7}
7	45.352	0.746	1.09157×10^{-7}	7.37083×10^{-7}	1.07517×10^{-7}	1.79872×10^{-7}
8	44.648	0.448	7.04502×10^{-8}	4.98284×10^{-7}	6.72217×10^{-8}	1.45749×10^{-7}

As in the previous sections, the in-band radiance $L_{zl}(\Lambda, \beta)$ is obtained by integrating $L_{zl}(\lambda, \Lambda, \beta)$ multiplied by the photometric passband transmission $S(\lambda)$.

The effect of the variation of the visual brightness of the zodiacal light with the heliocentric distance of the Earth r (in au) is modelled (see Leinert et al. 1980) by the factor $f_R = r^{-2.3}$.

Finally, the influence of the Earth position relative to the plane of the interplanetary dust cloud is also taken into account for high ecliptic latitudes ($|\beta| \geq 60^{\circ}$). It introduces a sinusoidal variation of ± 10 per cent in the zodiacal light brightness, having the most extreme values when the Earth is above or below this plane, and mean values at the nodes, placed at $\Omega=96^{\circ}$ (Leinert et al. 1998). This factor $f_{\rm S}$ takes the form

$$f_{\rm S}(\beta) = \begin{cases} 1 + 0.1 \sin(\Lambda_{\rm E} - \Omega) & \text{if } |\beta| \ge 60^{\circ}, \\ 1 & \text{otherwise,} \end{cases}$$
 (17)

where Λ_E is the ecliptic longitude of the Earth.

The in-band zodiacal radiance outside the Earth atmosphere at a point of ecliptic coordinates (Λ, β) can then be expressed by

$$L_{\rm zl}(\Lambda,\beta) = f_{\rm R} f_{\rm S}(\beta) L_{\rm zl_0}(\Lambda,\beta). \tag{18}$$

6 AIRGLOW

Airglow is a faint light emission originated in the upper atmosphere. It is caused by chemiluminescence, i.e. the emission of light (luminescence) from the decay of excited states of the products of a chemical reaction. In the case of the airglow, the chemiluminescence is triggered by the high-energy solar radiation. Airglow is emitted from several high-altitude atmosphere layers, starting around 90 km (mesopause), where the bright O I and OH emissions, together with fainter O_2 and Na D emissions, concentrate. Between ≈ 250 and \approx 300 km we found the emission of several O_I transitions. The airglow spectrum in the visible wavelengths is dominated by the OI green line at 558 nm (see Hart 2019a), the OI red lines at 630 and 636 nm (both produced in the layer between 200 and 300 km) and the FeO pseudo-continuum around 590 nm (Saran et al. 2011; Unterguggenberger et al. 2017). Finally, the hydrogen at geocorona contributes to the airglow in the H α line. Other lines like L α and L β are also generated in the geocorona by fluorescence, but they fall out our spectral range of interest. For a detailed analysis of the complex processes involved in the airglow generation, see Hart (2019a) or Noll et al. (2012).

Airglow is a highly variable source of the natural night sky brightness. The variability of the emissions due to the different constituents of the airglow (atomic and molecular O₂, Na, OH, ...) at different time scales (nightly, yearly, long term, ...) is thoroughly

analysed by Hart (2019a,b). For some of the main components it can reach 100 per cent (maximum minus minimum divided by the median). Also, airglow depends on the solar activity cycle (Patat 2008). Finally, the airglow emission could change with the geographic latitude, especially for lines originating in the mesopause (OH, O₂, Na_I D, FeO or most of the green O_I lines), and also with geomagnetic latitude, in particular for ionospheric lines such as the O_I lines at 630 and 636 nm. This makes the prediction of the actual value of the airglow emission difficult to model, and it could be considered as a free parameter if comparisons with real measurements are done. The basic model described here is only intended to provide a reference value and it shall be judiciously applied to any particular observation.

The most common model for the airglow brightness, i.e. the inband integrated airglow radiance $L_{\rm ag}(z)$ at different zenith angles z, is the one based on the factorization of the zenith brightness $L_{\rm ag}(0)$ and the zenith angle dependence given by the van Rhijn function (Leinert et al. 1998):

$$L_{\rm ag}(z) = \frac{L_{\rm ag}(0)}{\left(1 - \left[\frac{R}{R+H}\right]^2 \sin^2 z\right)^{\frac{1}{2}}},\tag{19}$$

where R = 6378 km is the Earth radius and H is the height above the Earth's surface of an assumed thin homogeneously emitting layer responsible for airglow. Different values have been quoted in the literature for H. Here we adopt H = 87 km (Hart 2019a). In general, the choice of H is only crucial at high zenith angles, but we need to bear in mind that, as mentioned above, different emissions are originating at different altitudes, and consequently, our choice of H could be inaccurate if strong emissions are originating at high altitude. For more accurate results, in particular in case of strong line emissions at different heights, a more complex airglow model, beyond the van Rhijn approximation with fixed H, should be used.

Some remarks should be made before applying this equation to our calculations. First, note that equation (19) is expressed in terms of in-band integrated radiances, $L_{\rm ag}(z)$, not of spectral radiances $L_{\rm ag}(\lambda,z)$, although it could be deemed approximately valid (with the same values for R and H) for $L_{\rm ag}(\lambda,z)$. Second, the van Rhijn formulation is meant to describe the at-the-layer radiance, $L_{\rm ag}(z)$, as seen at zenith angles z from the reference frame of an observer located at sea level (h=0 m). Strictly speaking, for observers located at heights h>0 m this equation should be slightly modified, to take into account that regions of the airglow layer seen at zenith angles z from sea level will appear to be at somewhat larger angles z'=z'(z,h), if seen by an observer at h>0 m. However, since for ground-based observers h is much smaller than R and H, this small correction can be ignored,

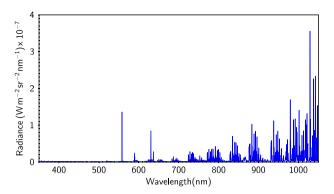


Figure 7. The airglow spectrum computed with the ESO's SKYCALC web interface for an altitude of 2640 m. Logarithm binning $\lambda/\Delta\lambda = 20\,000$ and Gaussian line spread function (LSF) convolution kernel [full width at half-maximum (FWHM) = 10 bins].

and equation (19) can be used to describe the at-the-layer radiance in their particular reference frames, irrespective from h.

The nadir-oriented spectral radiance at the layer, $L_{ag}(\lambda, 0)$, can be determined by observational measurements or from synthetic models. In the case of observational determinations, as in Hart (2019a), some other sky radiance components can contribute to some extent to the continuum value. For calculating the results presented in Section 8 of this work, we have used the Cerro Paranal Advanced Sky Model (Noll et al. 2012; Jones et al. 2013) to get a synthetic airglow line and continuum emission spectrum. The spectrum, computed from ESO's SKYCALC web interface, is calculated for the Cerro Paranal altitude (2640 m above the sea level) in the wavelength range from 350 to 1050 nm. For a reference spectrum, we set the value of the monthly averaged solar radio flux equal to 100 sfu $(1 \text{ sfu} = 10^{-22} \text{ W m}^{-2} \text{ s}^{-1})$, the approximate average value in the period 2009-2020 (one solar cycle) according to the data of the Canadian Space Weather Forecast Centre (CSWFC). This value can be set to other value for any specific application of the model. The outcoming spectrum is shown in Fig. 7. After correcting for the vertical atmospheric transmittance at Cerro Paranal Observatory, $T^{\text{CPO}}(\lambda, 0; h_0)$, also provided by the ESO's SKYCALC tool, we can determine the airglow radiance at the emission layer:

$$L_{\text{layer}}(\lambda, 0) = L_{\text{ag}}^{\text{CPO}}(\lambda, 0; h_0) / T^{\text{CPO}}(\lambda, 0; h_0). \tag{20}$$

As stated above, the final spectral airglow radiance is highly spatially and temporally variable and usually it should be considered as a free parameter. Its value could be set if more information about the airglow emission is available. This issue will be discussed in Section 8.1.

7 ATMOSPHERIC ATTENUATION AND SCATTERING

In the previous sections we have described the contributions of the astrophysical sources to the night sky brightness. Although slightly different models could be adopted for the background light or the zodiacal light, it is expected that the actual radiance outside the Earth's atmosphere does not differ significantly from the values given in this work. We have also discussed the contribution of the airglow, a highly variable source that should be considered as a free parameter of the model.

However, in order to obtain the local map of the night sky brightness, we need to deal with the effects of the Earth atmosphere on the outside radiance. In particular, the effect of the atmospheric attenuation and the scattering must to be added to the radiance coming from the astrophysical sources. Both effects are highly variable in time, since they strongly depend on the particular atmospheric conditions at the place and time of the observation. For the numerical examples presented in Section 8 of this work we adopted some standard values for the main atmospheric parameters. Different atmospheric states will give rise to different results.

The spectral radiance is modified by the terrestrial atmosphere by means of two interrelated processes. On the one hand, the radiance propagating toward the observer is attenuated by absorption and scattering by the atmospheric constituents. On the other hand, some amount of light initially propagating in different directions gets scattered into the observer's line of sight, being added to the recorded brightness. These two opposite effects stem from the same basic interaction processes at the atomic and molecular levels.

The attenuation of a beam along an atmospheric path of length d can be expressed in terms of the optical thickness $\tau(d)$ or, equivalently, of the transmittance T(d), as

$$L(d) = L(0) e^{-\tau(d)} = L(0) T(d), \tag{21}$$

where L(0) is the initial radiance and L(d) the radiance after travelling the distance d.

Henceforth we closely follow the formulation by Kocifaj (2007). Let us denote by $k_{\text{ext}}(h')$ the volume extinction coefficient at height h' above sea level, due to aerosols (A) and molecules (M), such that

$$k_{\text{ext}}(h') = k_{\text{ext}}^{\text{M}}(h') + k_{\text{ext}}^{\text{A}}(h')$$
 (22)

(which of course are wavelength dependent). For a layered atmosphere as implicitly assumed in equation (21), the optical thickness appearing in equation (22) can be computed as

$$\tau(d) = \tau(h_1, h_2) = \int_{h_1}^{h_2} k_{\text{ext}}(h') \, dh', \tag{23}$$

where dh' is a function of the trajectory followed by the light rays. Under the assumption of nearly rectilinear propagation between atmospheric layers, at an angle z with respect to the local zenith, we can write $dh' = M^M(z) dh$, and $dh' = M^A(z) dh$, where $M^M(z)$ and $M^A(z)$ are the molecular and aerosol airmasses, respectively, corresponding to z, and dh is measured along the vertical. Then,

$$\tau(h_1, h_2) = M^{M}(z) \int_{h_1}^{h_2} k_{\text{ext}}^{M}(h) \, dh + M^{A}(z) \int_{h_1}^{h_2} k_{\text{ext}}^{A}(h) \, dh.$$
 (24)

Under the assumption of exponential number density profiles for molecules and aerosols, the extinction coefficients take the form

$$k_{\text{ext}}^{\text{M}}(h) = \frac{\tau_0^{\text{M}}}{h_{\text{m}}} e^{-h/h_{\text{m}}}; \quad k_{\text{ext}}^{\text{A}}(h) = \frac{\tau_0^{\text{A}}}{h_{\text{a}}} e^{-h/h_{\text{a}}},$$
 (25)

where $\tau_0^{\rm M}$ and $\tau_0^{\rm A}$ are the (wavelength-dependent) molecular and aerosol optical thicknesses of the whole atmosphere, measured along a vertical path (z=0), for which $M^{\rm M}(0)=M^{\rm A}(0)=1$, and $h_{\rm m}$ and $h_{\rm a}$ are the molecular and aerosol scale heights, respectively. For our present calculations we have taken $h_{\rm m}=8$ km and $h_{\rm a}=1.54$ km.

The molecular (Rayleigh) component $\tau_0^M(\lambda)$ can be approximated analytically (Teillet 1990), with λ in μ m, by

$$\tau_0^{\mathcal{M}}(\lambda) = 0.00879 \,\lambda^{-4.09}.\tag{26}$$

And the aerosol component by

$$\tau_0^{\mathcal{A}}(\lambda) = \tau_0^{\mathcal{A}}(\lambda_0) \left(\frac{\lambda}{\lambda_0}\right)^{-\alpha},\tag{27}$$

 α being the Ångstrom exponent.

Data on atmospheric aerosols (aerosol optical depth and Ångstrom exponent) can be obtained from the Aerosol Robotic Network (AERONET) network (https://aeronet.gsfc.nasa.gov). For the results presented below, we chose $\tau_0^A(550 \text{ nm}) = 0.2$ and $\alpha = 1$. The choice of other values will affect the resulting night sky brightness, especially at high zenith angles, where the effect of atmospheric attenuation is more significant. For a highly accurate modelling, it is recommended to use the available local values.

The airmasses, at a first approximation, do not depend on wavelength. Then, the spectral attenuation across a path at zenith angle z, between the layers at h_1 and h_2 is

$$T(z, \lambda; h_1, h_2) = \exp \left\{ -M^{M}(z) \tau_0^{M}(\lambda) \left[e^{-h_1/h_m} - e^{-h_2/h_m} \right] -M^{A}(z) \tau_0^{A}(\lambda) \left[e^{-h_1/h_a} - e^{-h_2/h_a} \right] \right\}.$$
(28)

And the overall atmospheric attenuation, from $h_1 = h$ to $h_2 = \infty$

$$T(z, \lambda; h) = \exp \left\{ -M^{M}(z) \, \tau_{0}^{M}(\lambda) \, e^{-h/h_{m}} - M^{A}(z) \, \tau_{0}^{A}(\lambda) \, e^{-h/h_{a}} \right\}.$$
(29)

For the airmass we use the expression of Kasten & Young (1989), defined for all zenithal distances as

$$M(z) = \frac{1}{\cos(z) + 0.50572 (96.07995 - z)^{-1.6364}},$$
 (30)

with z in degrees.

Light from all other sky directions is also absorbed and scattered. When interacting with the atmospheric constituents located along the line of sight a fraction of this light gets scattered into the observer's field of view, adding to the detected radiance. Denoting by $L_0(\lambda, \alpha_s)$ the extra-atmospheric radiance from a differential patch of the sky of solid angle $d\omega$ around the generic direction α_s , the total radiance scattered into the observer line of sight $L_s(\lambda, \alpha, h)$ can be calculated as

$$L_{s}(\lambda, \boldsymbol{\alpha}, h) = \int_{\Omega} \Psi(\lambda, \boldsymbol{\alpha}, \boldsymbol{\alpha}_{s}, h) L_{0}(\lambda, \boldsymbol{\alpha}_{s}) d\omega, \tag{31}$$

where $\Psi(\lambda, \alpha, \alpha_s, h)$ is the function describing the spectral radiance scattered toward the observer along the direction α , per unit d ω , due to a unit amplitude radiant source located at α_s (all directions measured in the observer reference frame). The integral is extended to the whole hemisphere above the observer, Ω . For this work we have used the Kocifaj–Kránicz $\Psi(\lambda, \alpha, \alpha_s, h)$ described in equation (18) of Kocifaj & Kránicz (2011) with an effective scattering phase function composed of aerosol and molecular (Rayleigh) components weighted by the corresponding optical depths. The aerosol scattering phase function is described by a Henyey-Greenstein function with asymmetry parameter g = 0.90. The aerosol and molecular albedos have been set to 0.85 and 1, respectively. As indicated above regarding the aerosol optical depth, these particular values are used to provide the examples shown in Section 8. Better estimations of the natural night sky brightness for a given place and time could be obtained by using the appropriate local values.

The above formulation, adopted for the *Gaia–Hipparcos* map, requires performing an integration over the whole celestial hemisphere for every direction of observation, α . Several simpler but less accurate approaches have been proposed in the literature to account for the radiance scattered into the line of sight without calculating this integral. One of them is based on replacing $\tau_0^{M/A}(\lambda)$ in equation (29) by an *effective optical depth* $\tau_{0,\text{eff}}^{M/A}(\lambda) = \gamma \ \tau_0^{M/A}(\lambda)$ with $\gamma < 1$. The value of γ depends on the aerosol albedo and asymmetry parameter, and typical values are in the range 0.5–0.9 (Hong et al. 1998). This

is the approach used for diffuse sources in Duriscoe (2013), with $\gamma=0.75$, based on the empirical results of Kwon (1989). It can be a practical option for obtaining reasonably accurate results when computing time is a constraint, especially if the night sky brightness is evaluated in sky pixels of sufficient size to spatially average the contributions of the brightest stars.

8 RESULTS

In this section, we outline some applications of the GAMBONS model. The examples presented here have been calculated for particular choices of the atmospheric parameters (e.g. aerosol optical depth) and airglow zenith radiance. Quantitative comparisons with field measurements must take into account the actual state of the atmosphere and the airglow at the precise time of taking them.

All-sky maps like the ones described below can be computed for any location and time and freely downloaded from the GAMBONS website (https://gambons.fqa.ub.edu).

8.1 Comparison with all-sky images

As a first application, the GAMBONS model has been used to obtain all-sky maps of the brightness of the natural night sky. These maps provide a realistic description of what one can expect to observe in pristine locations free from light pollution sources, under different atmospheric conditions. The multiband capability of the GAMBONS approach allows to generate all-sky images in several bands.

The maps presented in this section were calculated for different photometric bands and are displayed in different units. In light pollution studies it is still frequent to work in the classical Johnson V band, reporting the brightness (i.e. the in-band radiances) in the logarithmic scale of magnitudes per square arcsecond. Let us remind that a region of the sky is said to have a brightness of m_V magnitudes per square arcsecond if each square arcsecond of its visual field gives rise – at the entrance pupil of the observing instrument – to the same irradiance a star of magnitude $m = m_V$ would produce (Bará et al. 2020). In this section, we use the V band as defined in Section 3.2. The magnitudes per square arcsecond are then calculated from the in-band radiances L (in W m⁻² sr⁻¹) as

$$m_V = -2.5 \log_{10}(L/L_{\rm r}),$$
 (32)

with $L_{\rm r} = 143.1685 \text{ W m}^{-2} \text{ sr}^{-1}$ as indicated in Table 2.

For human visual observations, the quantity of choice is the luminance, measured in SI units $cd\ m^{-2}$, corresponding to the photopic, mesopic, or scotopic adaptation states of the eye.

The all-sky images captured with some specific devices can be compared with the GAMBONS maps. One of these devices is the Sky Quality Camera (SQC) from Euromix Ltd. (Slovenia), a commercial Digital Single-Lens Reflex (DSLR) camera with fish-eye lens to evaluate the night sky brightness in the whole sky. This kind of devices is a versatile solution to measure the night sky brightness (see Hänel et al. 2018), providing directional information. The SQC software allows processing raw all-sky images and provides different products, including calibrated data in the V band (Jechow et al. 2018; Vandersteen et al. 2020). In Fig. 8, we show the comparison of two of those images with our model. The images were taken by one of the authors (SJR) from Senyús ($\lambda = 1^{\circ}12^{'}18^{''}$ E, $\phi = 42^{\circ}13^{'}52^{''}$ N, h = 1142 m), a pristine dark site with very low light pollution near the Montsec protected area (Catalonia, Spain) at two different epochs. The agreement between the SQC data and the model is fairly good. The main features of the sky (zodiacal light, clearly visible in the

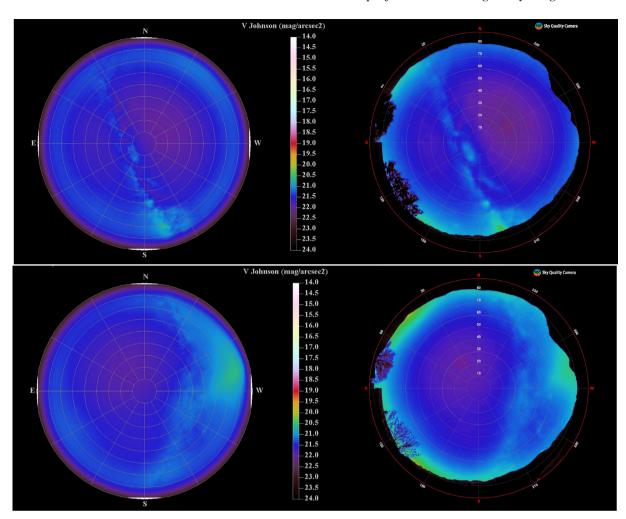


Figure 8. Comparison of GAMBONS maps (left) and SQC images (right) in V Johnson passband for Senyús ($\lambda = 1^{\circ}12^{'}18^{''}$ E, $\phi = 42^{\circ}13^{'}52^{''}$ N, h = 1142 m, Catalonia, Spain). The colour scale represents the sky brightness in Johnson V mag arcsec⁻². Bright patches near the horizon in the SQC images are due to light pollution. Top: 2018 August 10, 21:46 UT. Bottom: 2019 March 23, 20:28 UT.

image of the bottom panel in the west direction, or Milky way) in these SQC images are coincident with those predicted by GAMBONS.

Note that this is a qualitative comparison. As mentioned in Section 6, the radiance of the airglow is very variable both in short and large time-scales, and for many practical applications it may be considered as a free parameter. We have used the value of the solar radio flux from the CSWFC to get the airglow spectra for the both epochs, but it could not account for local or very short variations of the airglow radiance. The same applies to the atmospheric parameters related with the attenuation. After some tests, the best match is obtained with $\tau_0^{\rm A}=0.15$ and $\alpha=1$. Furthermore, SQC images record light pollution, and its V band could not coincide with the V Johnson band used in this work, as it is derived from the RGB image taken with a DSLR camera. This could introduce some differences between SQC and GAMBONS values, function of the colour of the sky.

The GAMBONS model can also be used to determine the amount of artificial light polluting the night sky. As in the previous case, it requires the fitting of the parameters not available from direct measurements, as e.g. the airglow, and in some cases the aerosol properties, if they are unknown. The fit could be laborious due to the complex spatial features of the actual airglow and attenuation. The model can be easily extended to other photometric bands used

in light pollution studies, as e.g. the Sky Quality Meter (SQM) and Telescope Encoder and Sky Sensor Wifi (TESS-W; Bará, Tapia & Zamorano 2019) or RGB (Sánchez de Miguel et al. 2019; Kolláth et al. 2020). Both topics will be addressed in forthcoming papers. By way of example, in Fig. 9 we show the visual scotopic and photopic all sky maps for the same location and date. They are expressed in luminance units μ cd m⁻².

8.2 Relative contributions to the night sky brightness

The model allows to compute the expected contribution of the different sources (ISL, DGL + ELB, zodiacal light, and airglow) to the total night sky brightness (moonless and cloudless) under different conditions (i.e. different locations, times, or airglow intensities). The contribution also depends on the extent of the field of view that we consider. We refer here as *zenith values* to the average radiance in a circular region of radius 10° around the zenith with no weights applied to the different points.

We have run the model for a mid-latitude location ($\phi=40^\circ$) and the atmospheric conditions given in Section 7 and the reference (constant) airglow spectrum described in Section 6, and averaged the contribution of the different sources over a year. The model was run

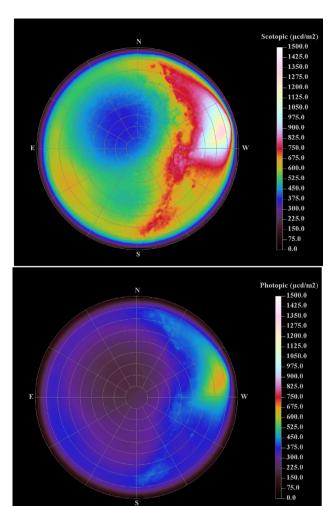


Figure 9. All sky luminance map (μ cd m⁻²) for observers with scotopic (top) and photopic (bottom) luminance adaptation. 2019 March 23, 20:28 UT, Senyús ($\lambda = 1^{\circ}12^{'}18^{''}$ E, $\phi = 42^{\circ}13^{'}52^{''}$ N, h = 1142 m, Catalonia, Spain).

Table 4. Mean contributions to zenith natural night sky brightness for a latitude equal to 40° N.

Component	Mean radiances (nW m ⁻² sr ⁻¹)	Percentage (%)
Airglow	165	47.7
Zodiacal light	95	27.5
Integrated starlight	75	21.7
Diffuse galactic light	10	2.9
Extragalactic background light	0.8	≈0.2

in 1 h intervals during the astronomical night (i.e. Sun more than 18° below the horizon). The result for the V band is shown in Table 4. Note that these values are sensitive to the airglow intensity, which is highly variable.

The contributions also depend on the observer latitude, as shown in Fig. 10. Zodiacal light has a larger impact in locations close to the Equator when we observe near the zenith, as the plane of the ecliptic reaches greater heights above the horizon at these latitudes.

Finally, the radiances are also a function of the observation time. The presence or absence of the Milky Way and the height of the ecliptic above the horizon are the two main factors that modulate the relative contributions to the sky brightness throughout the year.

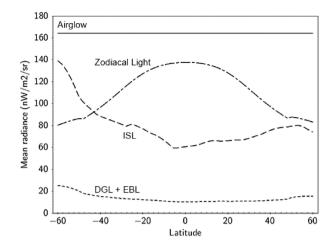


Figure 10. Annual average radiance at midnight in the Johnson V band at zenith for the different contributors to the natural sky brightness as function of the observer latitude.

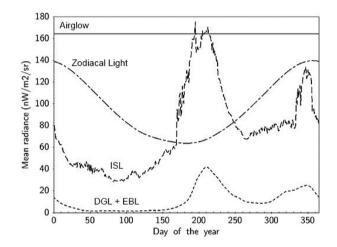


Figure 11. Midnight radiance at zenith in the Johnson V band for the different contributors to the natural sky brightness as function of the day of the year for an observer at latitude equal to 40° N.

As shown in Fig. 11, for the midnight radiance around the zenith for an observer at mid-latitudes, the Milky Way could become the brightness component. As it can be observed in both figures, we assume a constant airglow radiance, regardless of the location and time.

8.3 Seasonal variation of the zenith sky brightness

For a given location, the natural night sky brightness is highly dependent on the observing time. As we have already mentioned, the presence or absence of the Milky Way or the altitude of the ecliptic plane above the horizon strongly determine the brightness of the sky. We must bear this in mind when trying to characterize the darkness of a given place. According to our model, for zenith measurements, the natural variation of the night sky brightness along the year due to the variation of the astrophysical contributions can reach more than 0.6 mag. As an example, Fig. 12 shows the variation of the V magnitude around the zenith (i.e. a 10° circular region around it) for an observer at $\phi = 40^\circ$ latitude and h = 1000 m.a.s.l. The figure plots only the times when the height of the Sun above the horizon is $<-18^\circ$. The presence of the Milky Way in two different epochs

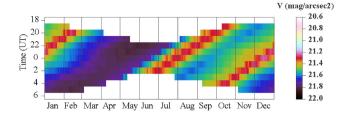


Figure 12. Variation of the natural night sky brightness (mag_V arcsec⁻² at zenith) along the year for an observer at $\phi = 40^{\circ}$ and h = 1000 m.a.s.l. Note that the colour scale is zoomed with regard to Fig. 8.

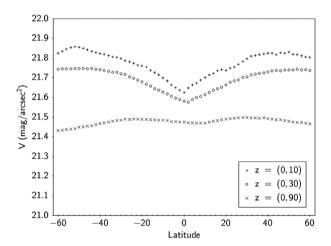


Figure 13. Darkest value of the natural night sky brightness (in $mag_V \operatorname{arcsec}^{-2}$) as a function of the latitude for three different fields of view: 10° around the zenith (zenith angle z=(0,10)); 30° around the zenith (z=(0,30)); and all-sky (z=(0,90)). The atmospheric and airglow conditions are the ones assumed in this section.

of the year is clearly visible. The zodiacal light is also modulated by the position of the Earth above the ecliptic plane and the Sun–Earth distance. The variation is smoother if a wider area of the sky is considered. But even for relatively large fields of view of tens of degrees, the variation of the natural night sky brightness is up to several tenths of $mag_V arcsec^{-2}$. Again, we consider constant airglow.

8.4 How dark are the darkest natural night skies?

There might be a question about the darkest value the natural night sky can attain, in absence of any source of artificial light pollution. This value is significant for light pollution studies, as it establishes an upper limit to the actual measurements of the sky darkness. Darker measurements could indicate higher atmospheric attenuation or sporadic lower airglow radiances, not necessarily being indicative of less polluted ('darker') skies. The darkest value of the zenith natural night sky depends on the observer latitude and the time of observation, as well as on the atmospheric conditions, altitude of the observer above sea level, and airglow intensity. Because of attenuation, less stars are expected to be perceived from sites at lower altitudes above sea level, as well as at higher zenith angles, as recently analysed in Cinzano & Falchi (2020). For a given latitude, using our model (with the particular airglow and atmospheric conditions assumed in this section), we are able to determine the minimum value of the night sky brightness (i.e. the darkest sky at a given latitude). It is shown in Fig. 13. The darkest skies (averaged over 10° around the zenith) are found at mid-latitudes in both hemispheres.

The presence of the ecliptic plane near the zenith for observers near the Equator prevents them to reach the darkest values. This effect almost disappears if we considered the full sky dome, but it is well visible even for a wide field of view $(0^{\circ} < z < 30^{\circ})$.

9 CONCLUSIONS

We present the GAMBONS model of the natural night sky brightness. The model considers each of the contributors to the sky brightness (i.e. the ISL, the DGL, the EBL, the zodiacal light, and the airglow), plus the atmospheric attenuation and scattering into the field of view. The main novelty with respect to previous models is the use of the Gaia-DR2 catalogue to evaluate the ISL. Furthermore, the photometric data in Gaia-DR2 (G, $G_{\rm BP}$, and $G_{\rm RP}$ bands) allow to transform from Gaia photon fluxes to radiances in any other photometric band. In this way, GAMBONS becomes a multiband model to study the natural night sky brightness.

In particular, although it amounts less than the 3 per cent of the sky brightness, we have put an effort to evaluate the contribution of the astrophysical diffuse light (DGL plus EBL), by using the relation between the emission at 100 μm and the optical emission. The spectrum template of the airglow used by GAMBONS is based on the Cerro Paranal Advanced Sky Model, through the ESO SKYCALC tool.

Two highly variable inputs to the model are the absolute zenith value of the airglow radiance and the aerosol properties at the time of observation. They vary from place to place and in different time-scales. Information on aerosol properties at a large network of monitoring sites worldwide can be obtained from AERONET.

GAMBONS is expected to enable multiple applications in the study and characterization of the natural night sky brightness. We have pointed out several of them. In particular we have shown the variation of the sky darkness as function of the latitude of the observer and time, as well as the variation of the contribution of the different sources of natural light with these two variables. Furthermore, GAMBONS could help, if reliable determinations of airglow intensity and atmospheric conditions are available, to establish a reference value of the sky darkness, for any given location in a moonless and cloudless night, against which to evaluate the measurements reported by the observers. In a forthcoming paper we will deal with the use of GAMBONS to remove the natural night sky brightness from the raw all-sky images in order to quantitatively estimate the levels of light pollution, and to put to the test the predictions obtained from models of atmospheric propagation of artificial light at night.

GAMBONS and the data related are accessible via web at https://gambons.fqa.ub.edu.

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DATA AVAILABILITY

The data underlying this paper are available in the paper and in its online supplementary material.

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SUPPORTING INFORMATION

Supplementary data are available at MNRAS online.

Table 3. Radiances (W m⁻² sr⁻¹) outside the Earth atmosphere.

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